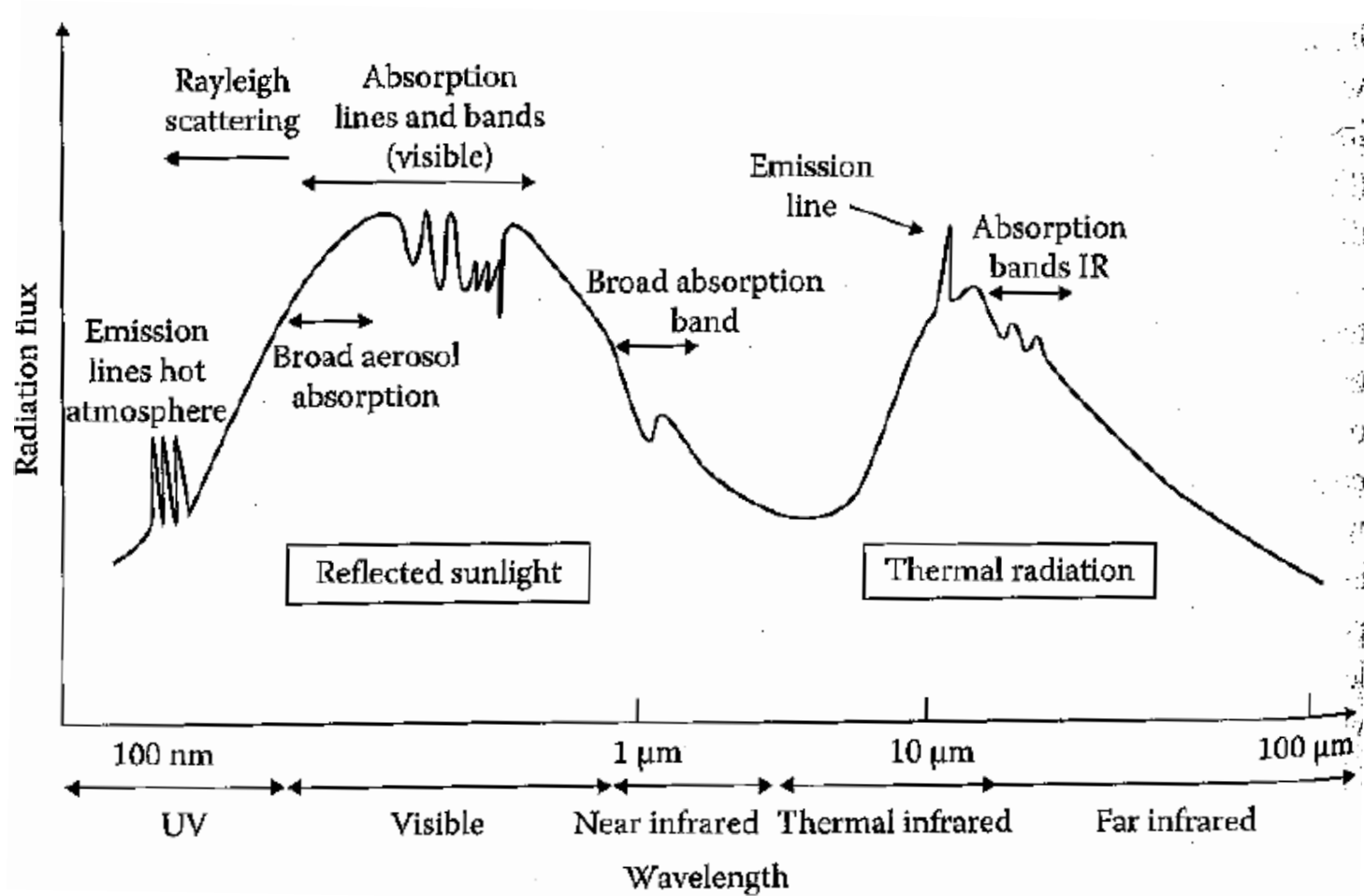


Abundances in the solar system: what we achieved and how much was it dependent on space probes?

Sandrine VINATIER
LESIA, Observatoire de Paris-Meudon

Introduction: a typical planet radiation



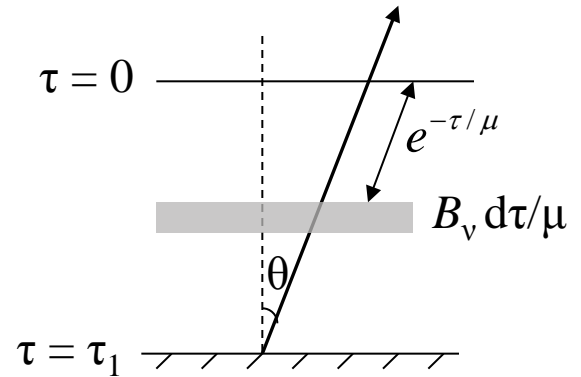
Introduction: thermal IR emission

■ Outgoing thermal emission of a planet

$$I_\nu(\mu, 0) = \underbrace{B_\nu(T_1)}_{\text{Surface/clouds}} e^{-\tau_1/\mu} + \frac{1}{\mu} \int_0^{\tau_1} \underbrace{B_\nu(T)}_{\text{atmosphere}} e^{-\tau/\mu} d\tau$$

Planck function,
depends on T

opacity: depends on composition
(molecular abundances)



Pb: How to decouple temperature and composition contributions?

■ Temperature:

- fit of the observed I_ν
- a gas uniformly mixed with known abundance

} $\Rightarrow T$

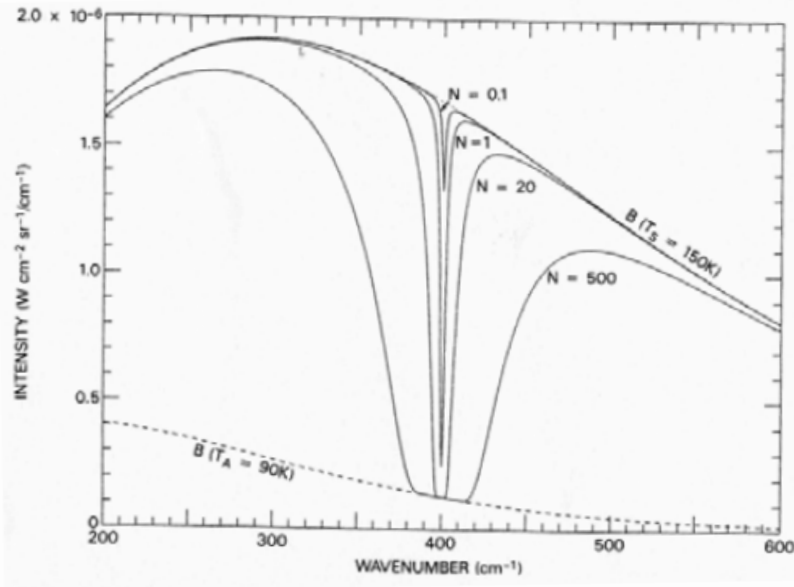
■ Abundance:

- fit of the observed I_ν
- T is known

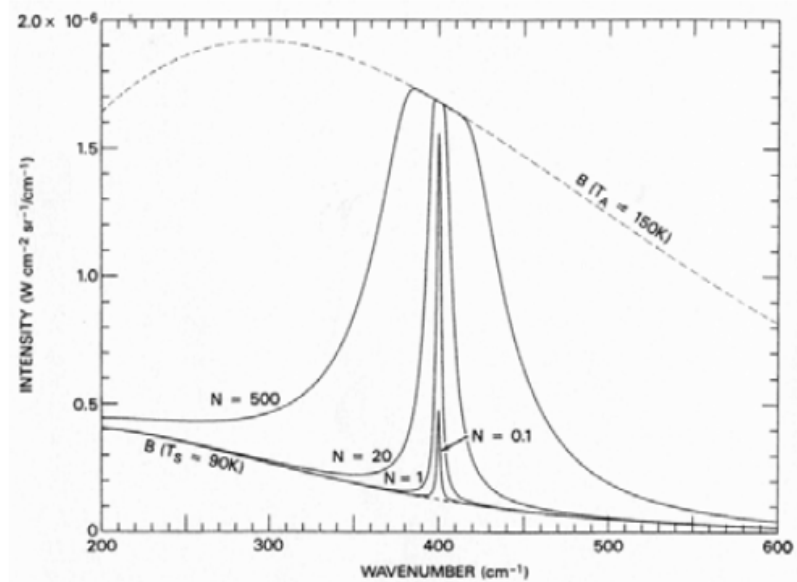
} $\Rightarrow \tau \Rightarrow q$

Introduction: radiative transfer

Absorption line



Emission line



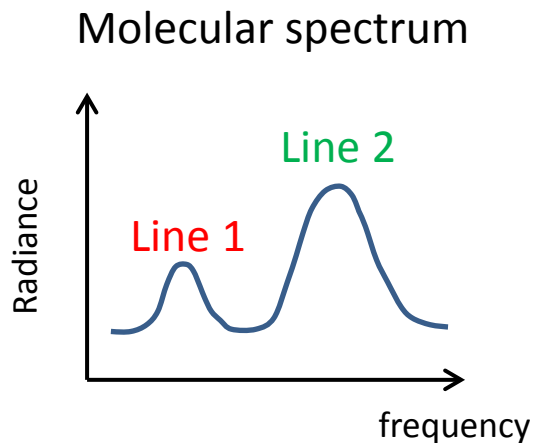
Width of the line depends on the gas density: **the higher is the abundance, the larger is the line** ⇔ **pressure broadening**

Absorption/emission depends on the temperature gradient in the region of the radiation emission (where $\tau = 1$).

Introduction: line opacity and level probed

■ How to probe several levels?

For a given molecule, lines of different intensities originate from different levels



Line 1 has a small absorption coefficient => probe deeper in the atmosphere.

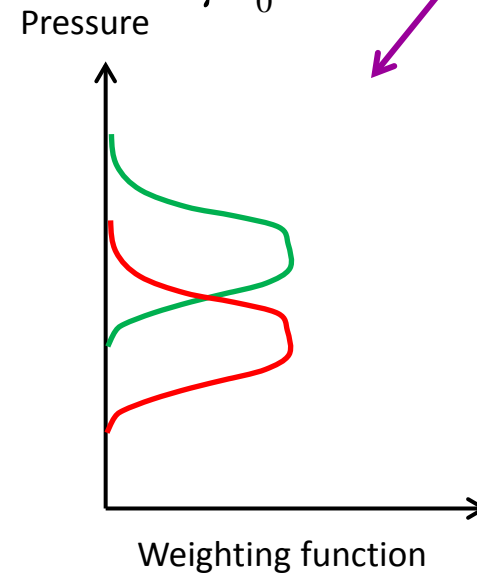
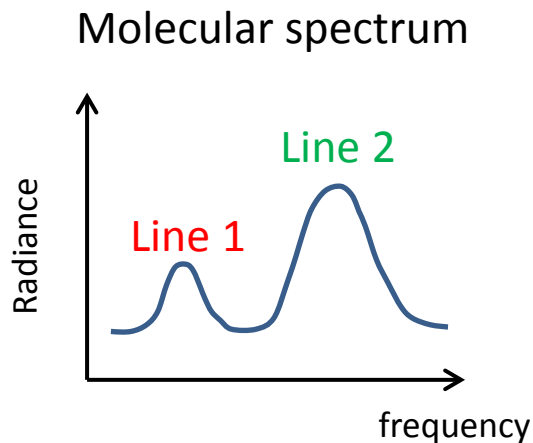
Line 2 has a high absorption coefficient => probe higher in the atmosphere.

Introduction: line opacity and level probed

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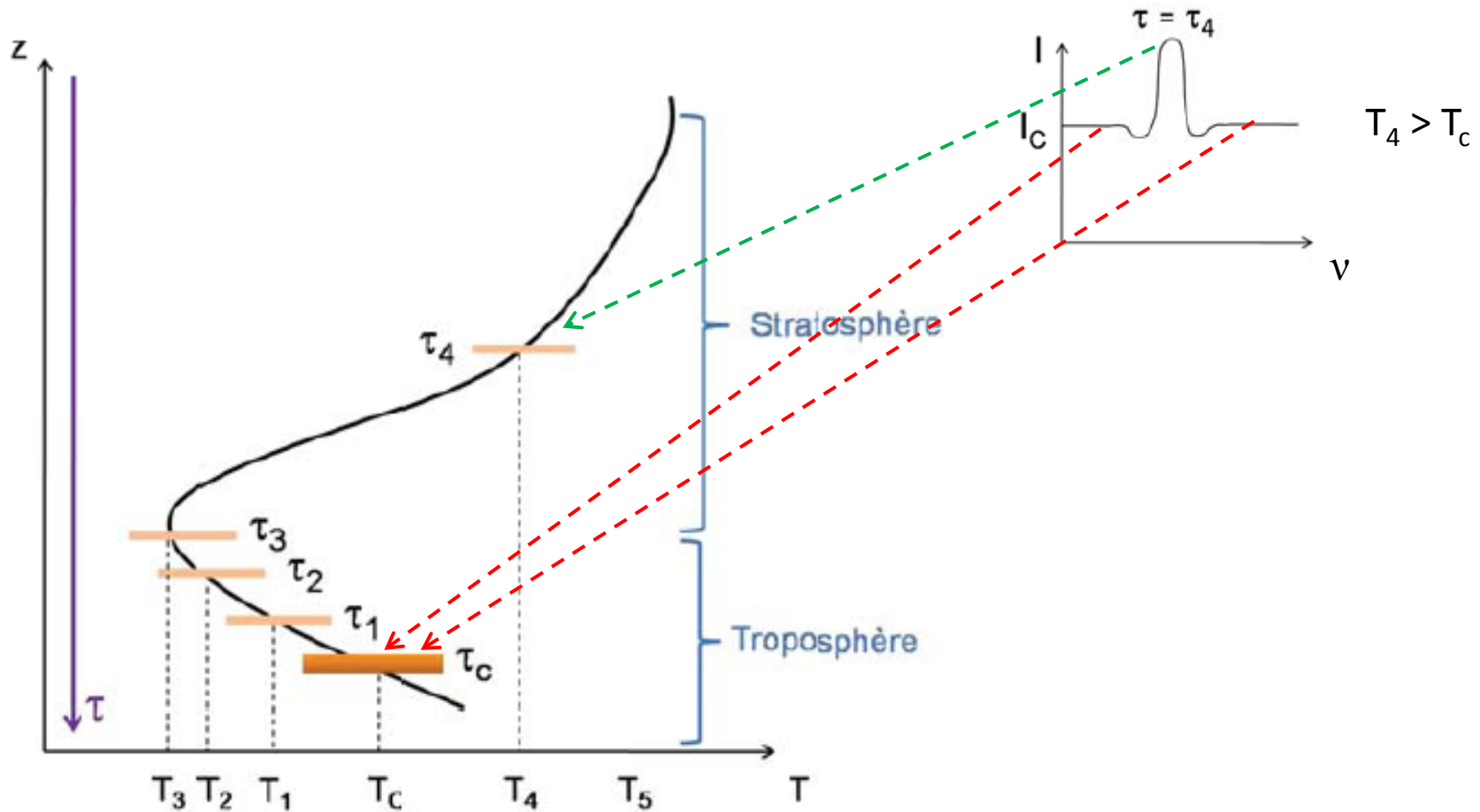
$$I_\nu(\mu, 0) = B_\nu(T_1)e^{-\tau_1/\mu} + \frac{1}{\mu} \int_0^\infty B_\nu(p) WF(p) d \ln p$$



Line 1 has a small absorption coefficient => probe deeper in the atmosphere.

Line 2 has a high absorption coefficient => probe higher in the atmosphere.

Introduction: line profile vs thermal profile

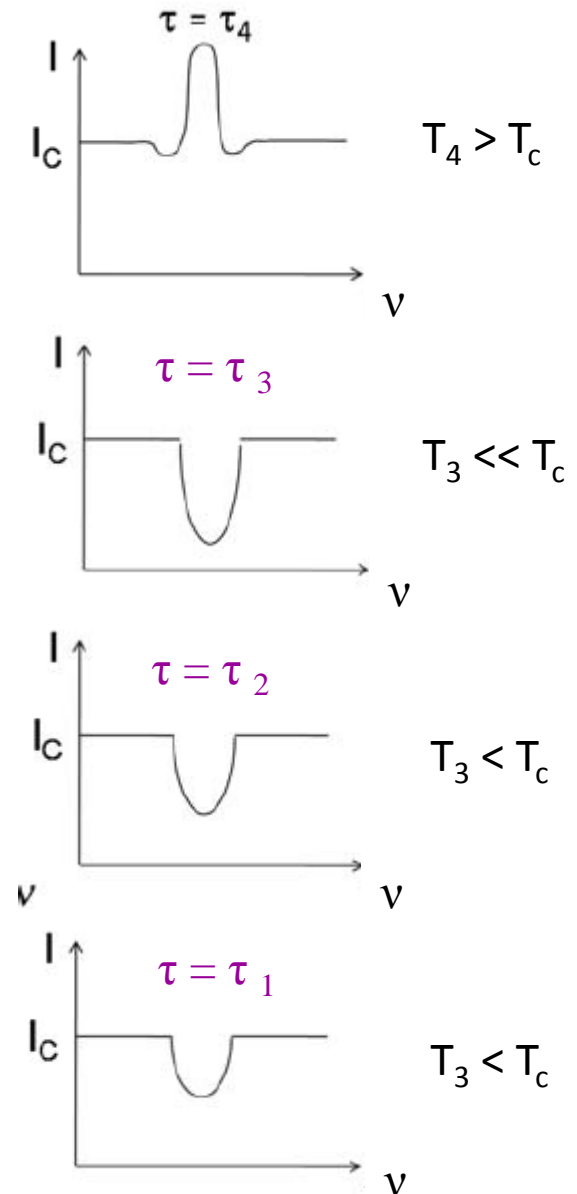
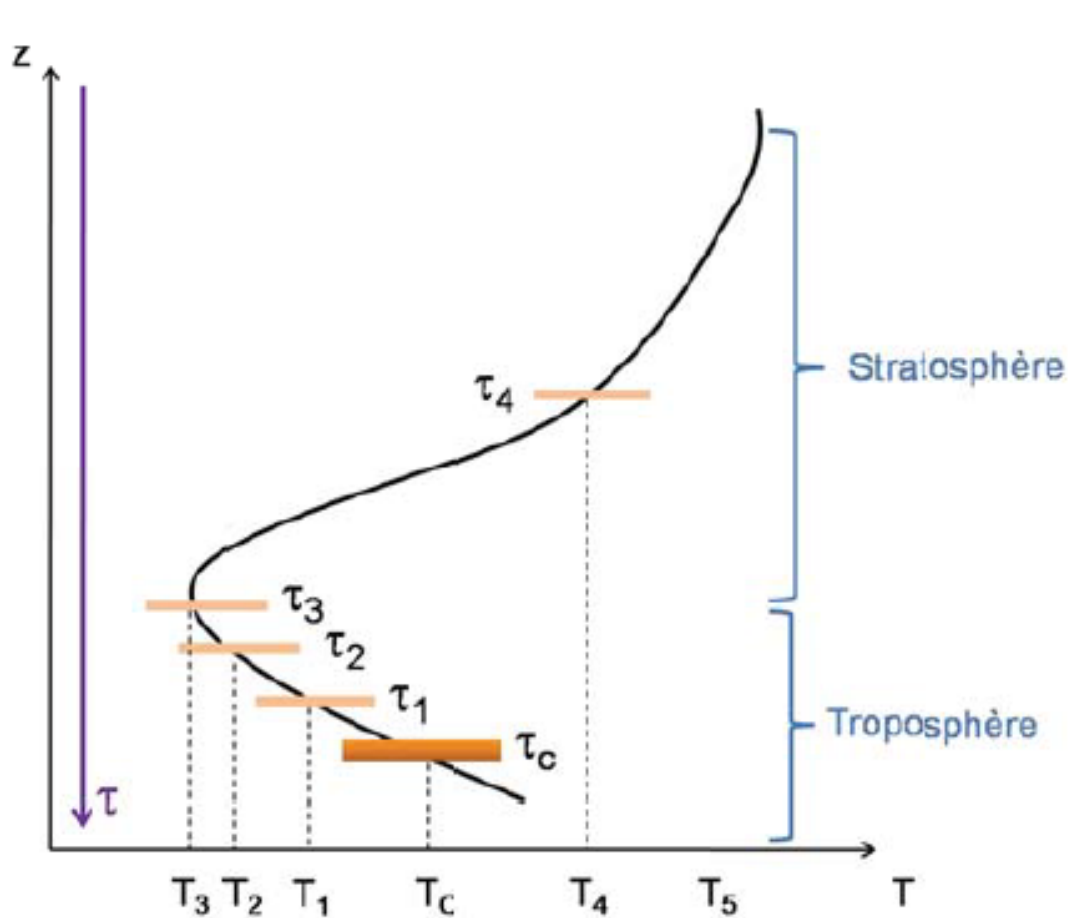


Region of positive thermal gradient => **emission line**

Region of negative thermal gradient => **absorption line**

If isothermal profile => **no line**

Introduction: line profile vs thermal profile



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 Region of negative thermal gradient => **absorption line**
 If isothermal profile => **no line**

Titan

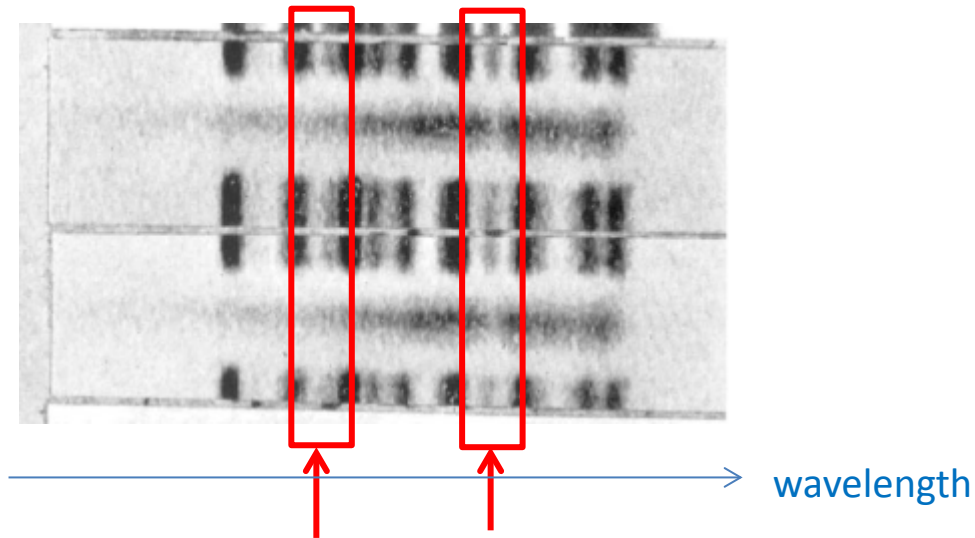
Titan's composition: an historical overview

- 1925: Jeans studied the atmospheric escape process => Titan should have kept an atmosphere.
 - gas of molecular weight ≥ 16 : possibly Ar, Ne, N₂, CH₄, NH₃ from thermodynamical theory
 - no H₂ and He because they are too light and should have escaped

Titan's composition: an historical overview

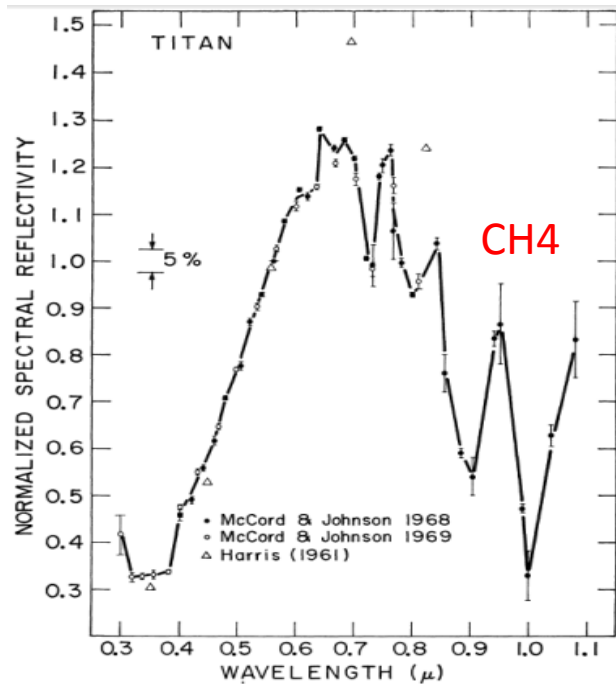
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- 1944 : first detection of CH₄ (Kuiper, 1944)



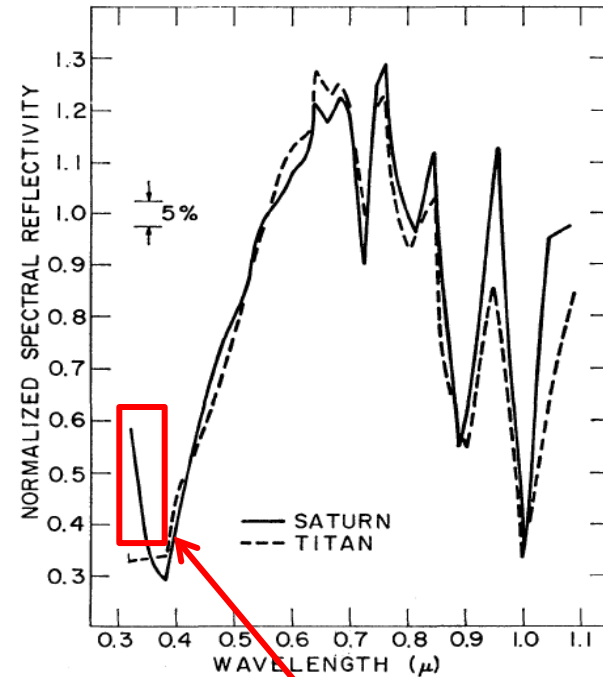
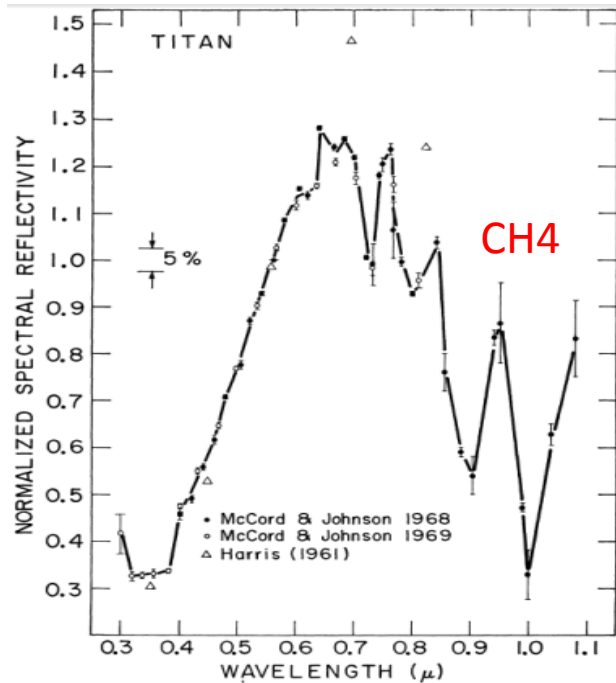
Titan's spectroscopy: reflected light

- 1961 – 1973 : UV ground-based observations (McCord et al. 1971) and space observations with the Orbiting Astronomical Observatory (Caldwell et al., 1973)



Titan's spectroscopy: reflected light

- 1961 – 1973 : UV ground-based observations (McCord et al. 1971) and space observations with the Orbiting Astronomical Observatory (Caldwell et al., 1973)



Rayleigh scattering

- Very similar spectra for Titan and Saturn except in UV => 2 different models for Titan
- low density atmosphere => Rayleigh scattering is not efficient.
 - Rayleigh scattering exists but is masked by an UV absorber = haze.
- If an UV absorber exists => deposition of energy at high altitude => T \uparrow with height

Titan's spectroscopy: reflected light

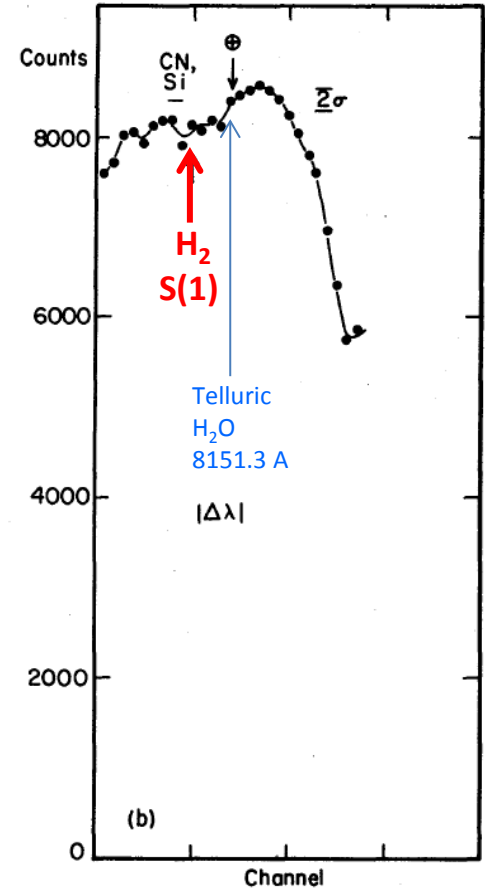
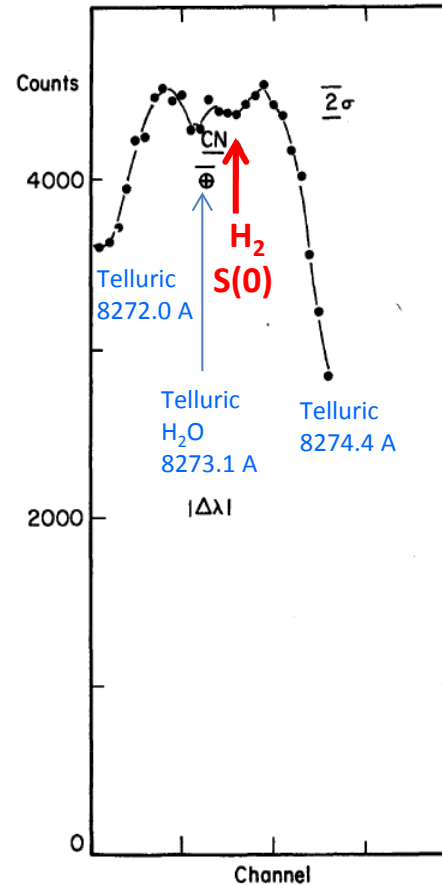
■ Possible detection of H₂: unexpected from Jean's escape calculations

1972 : possible detection of H₂ by Trafton et al.

McDonald Observatory (2 m)

1975 : confirmation of the detection of the S(1) band by Trafton et al.

=> presence of another gas that would inhibit H₂ escape.



In 1972 : Titan's atmosphere composition : CH₄, H₂, aerosols + ???

Titan's spectroscopy: reflected light

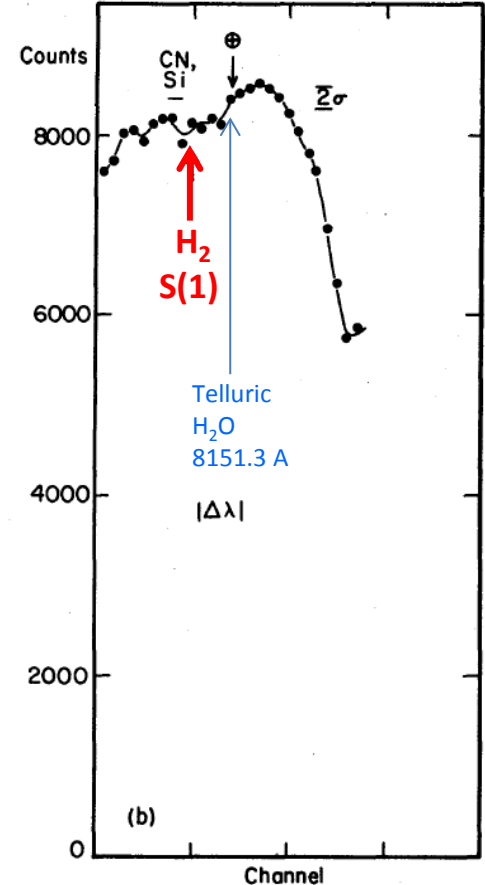
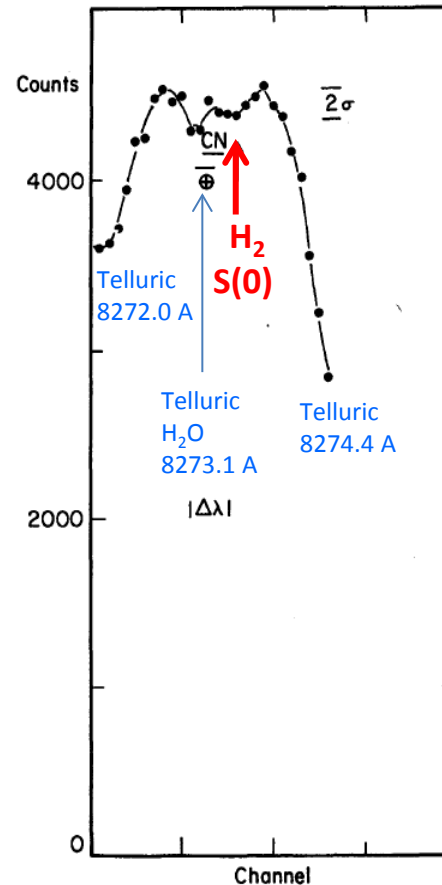
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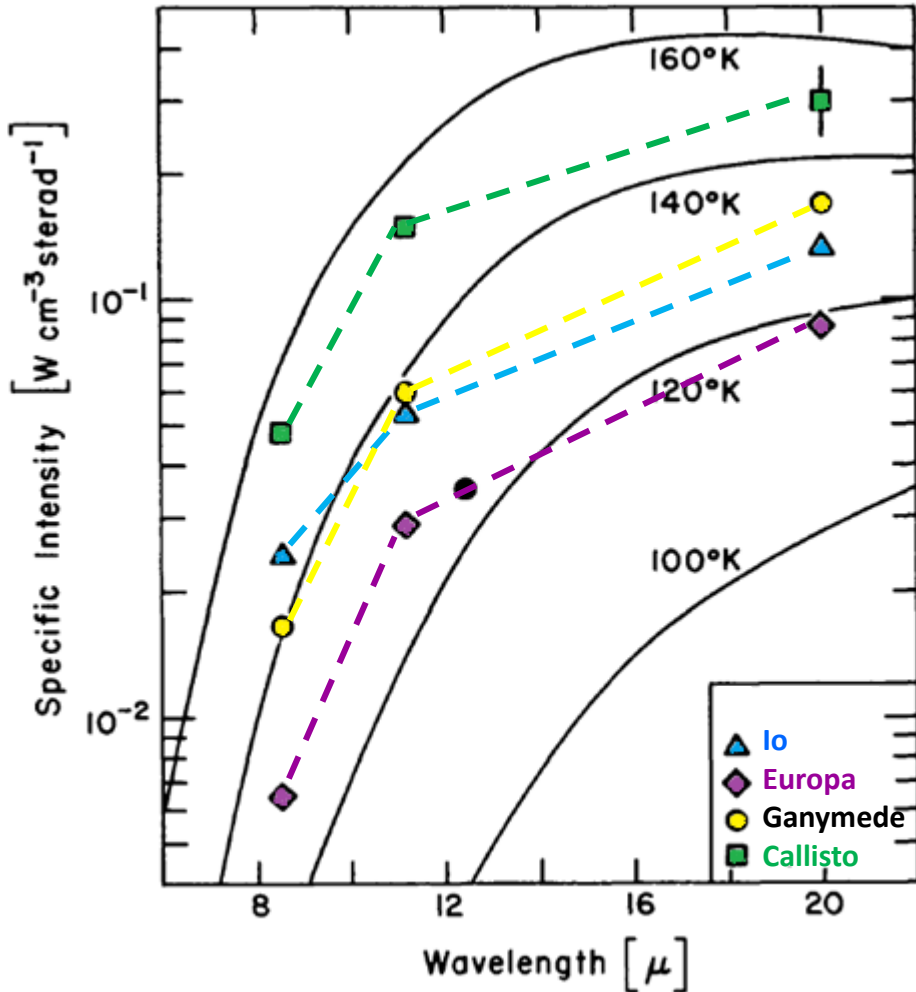
=> presence of another gas that would inhibit H₂ escape.



In 1972 : Titan's atmosphere composition : CH₄, H₂, aerosols + ???

Rmq: In 1981, Trafton et al. mentioned that their 1972 and 1975 H₂ observations were mostly due to CH₄.

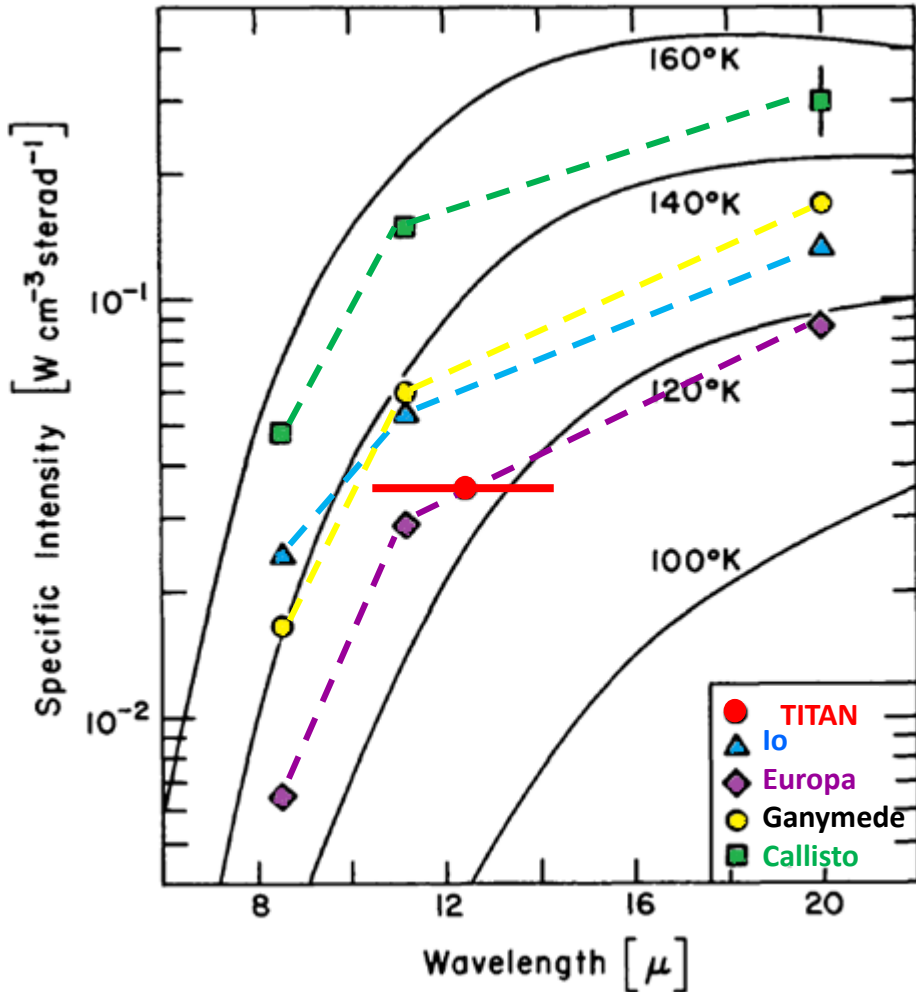
Titan's photometry: thermal emission (12 - 20 μm)



Jupiter satellites (no atmosphere)
temperatures : close to BB emission (Gillet et al., 1970).

Titan : farther from the Sun, $A_b = 0.29$
 \Rightarrow Titan's equilibrium temperature should be 82 K.

Titan's photometry: thermal emission (12 - 20 μm)

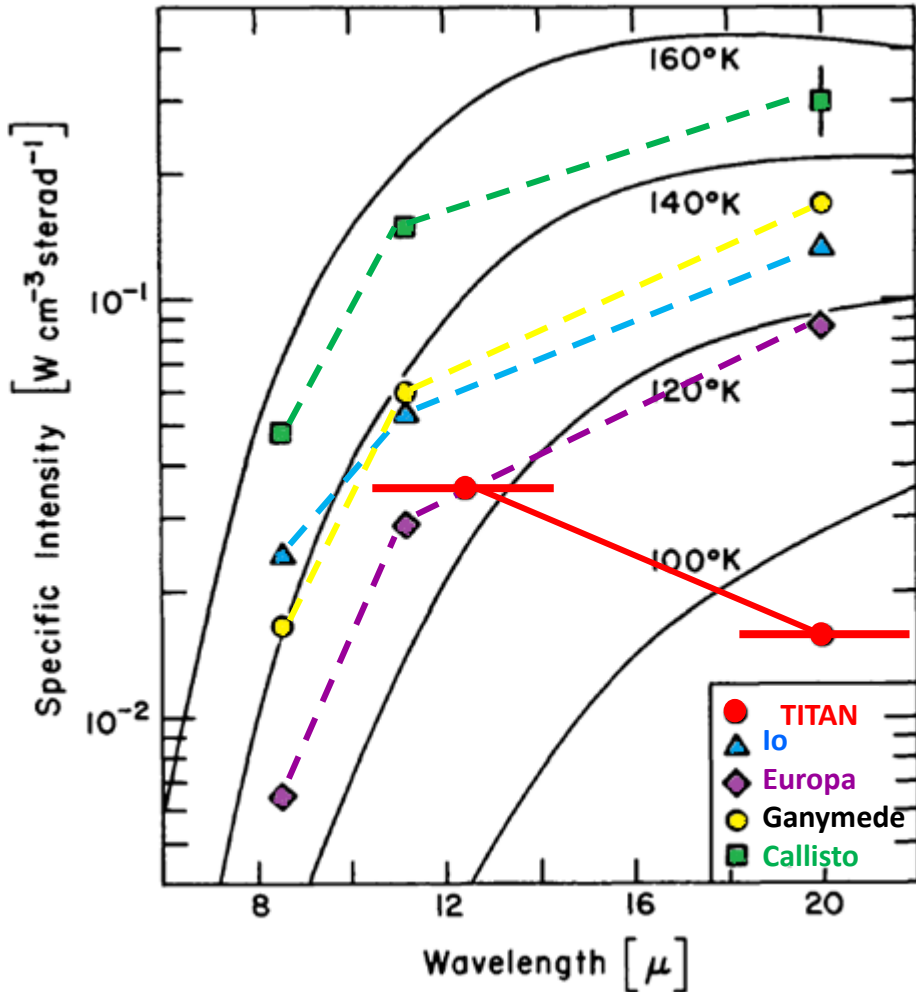


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Titan at 12 μm (Allen et al., 1971) :
used O'Brien Observatory (76 cm)
 $\rightarrow T_b = 125 \pm 5 \text{ K}$ (hyp: $R_{\text{Titan}} = 2405 \text{ km}$)
 \Rightarrow Greenhouse effect in a thin atmosphere

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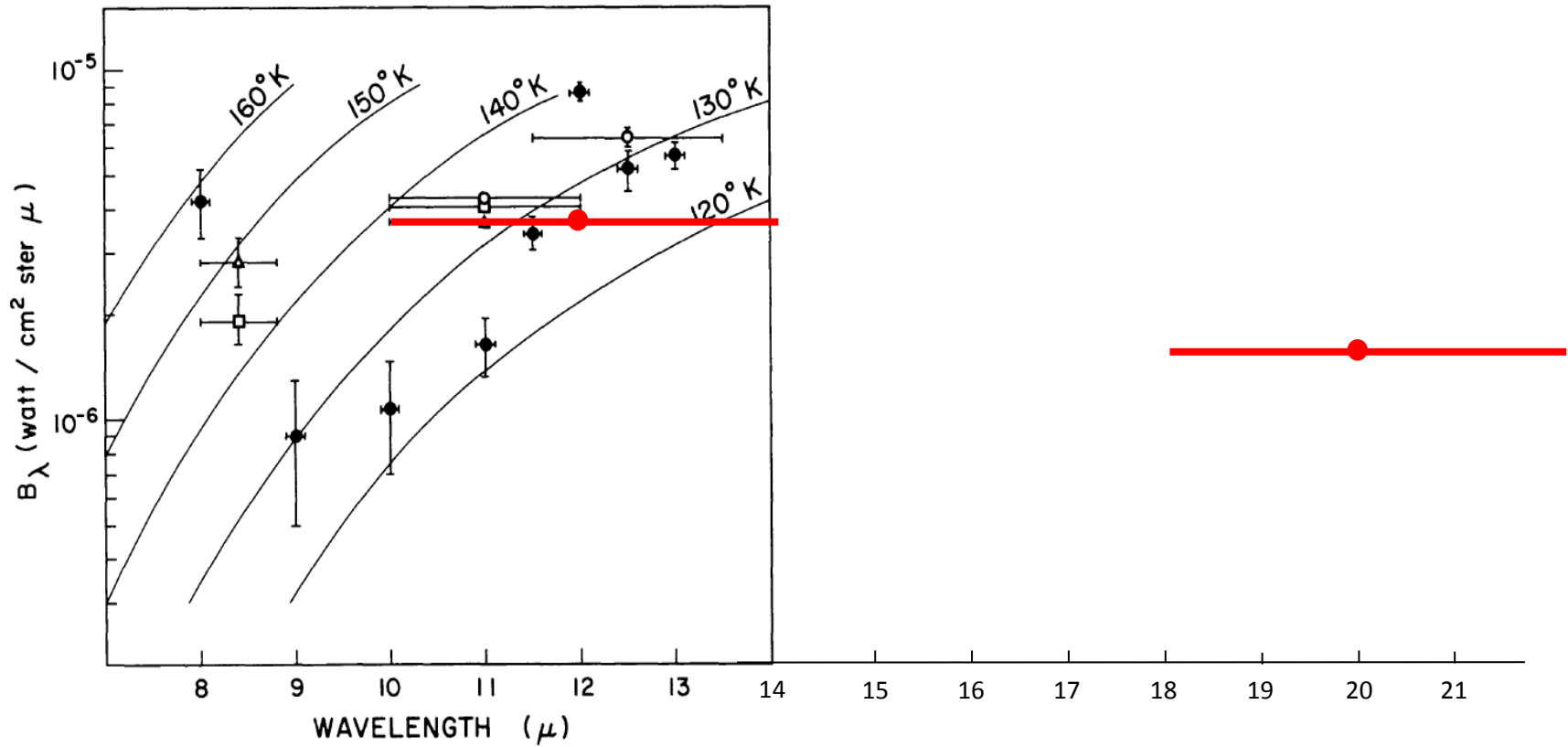
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Titan at 20 μm (Morrison et al., 1972) :
Mauna Kea observatory (2.24 m).
interesting because = peak of the BB emission
 $\rightarrow T_b = 93$ K, **puzzling!** ($R_{\text{Titan}} = 2550$ km)

2 possibilities : - surface emissivity is very low

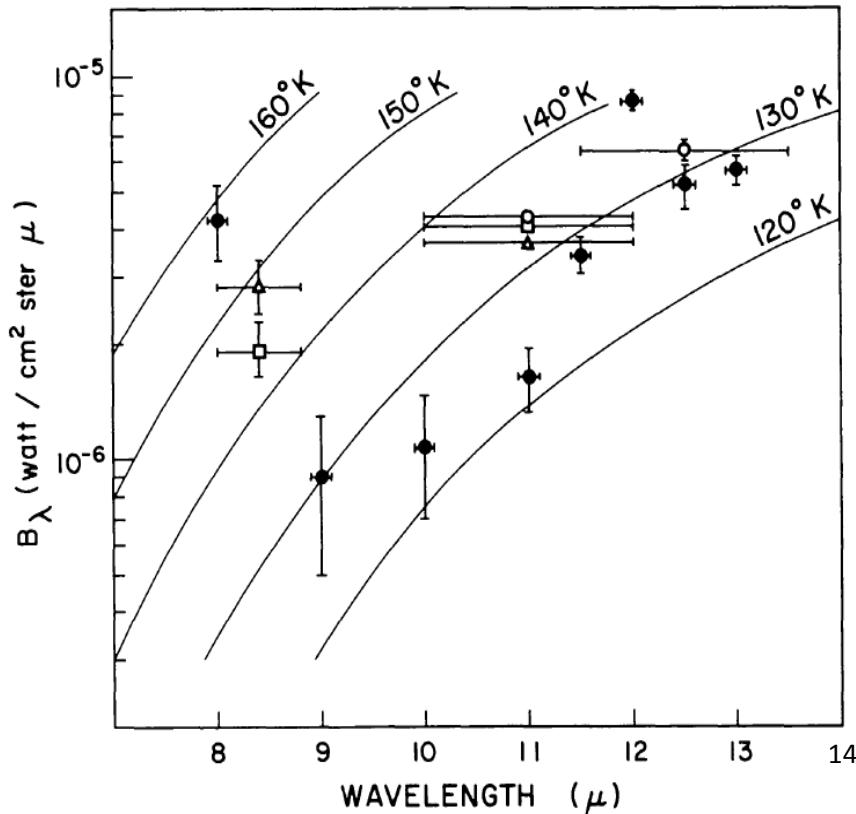
- opacity of the atmosphere very high at 20 μm \rightarrow possibly H_2 (collision induced), then the surface pressure should be ~ 1 atm

Titan's spectroscopy: thermal emission (8 - 13 μm)



Gillet et al., (1973) : observations at Mount Lemmon (152 cm)

Titan's spectroscopy: thermal emission (8 - 13 μm)



Gillet et al., (1973) : observations at Mount Lemmon (152 cm), assumed : $R_{\text{Titan}} = 2440$ km.

2 possible interpretations :

- If $T \searrow$ with height \Rightarrow strong absorber around $10 \mu\text{m}$: NH_3 ($10.5 \mu\text{m}$) ?

- If $T \nearrow$ with height \Rightarrow strong emission features in the $8 \mu\text{m}$ and $12.5 \mu\text{m}$ regions : CH_4 ($7.7 \mu\text{m}$), C_2H_6 ($12.2 \mu\text{m}$)

Need of a better spectral resolution

Titan's atmosphere: the first models

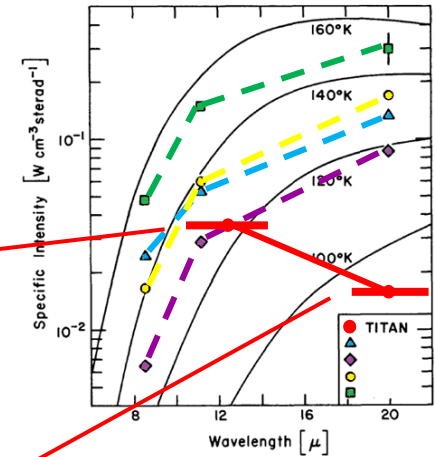
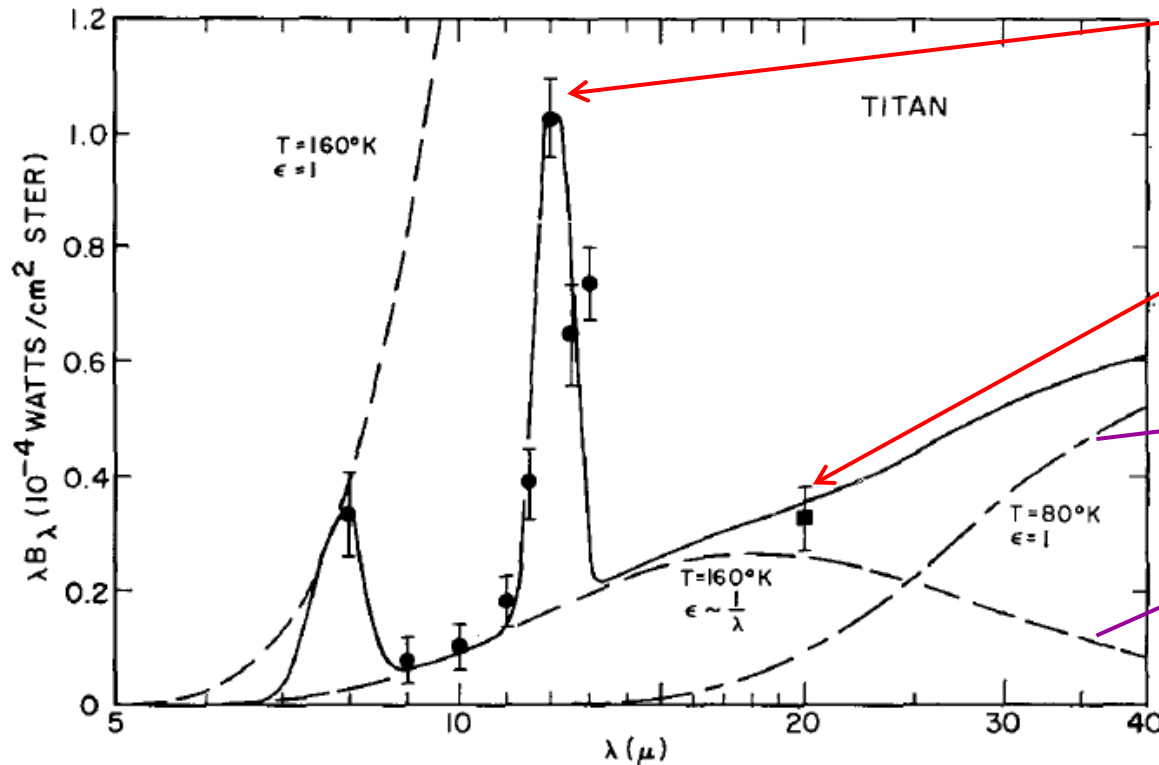
- Greenhouse effect in a massive atmosphere (Pollack, 1973)
- Strong atmospheric inversion in a much less massive atmosphere (Danielson et al., 1973)
- Greenhouse effect + inversion (Low and Rieke, 1974)

- Photochemical models (Strobel, 1974)

Titan: confirmation of a temperature inversion

Danielson 1973: model including:

- a temperature inversion
- Opacities of CH_4 , C_2H_6 , and aerosols

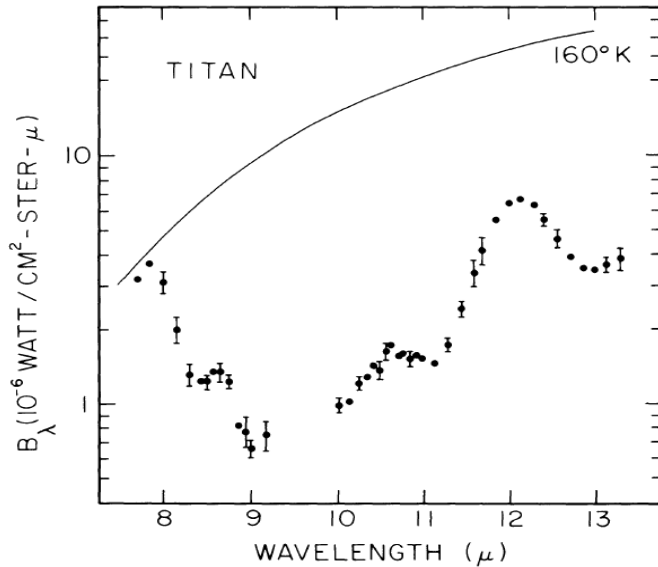


Radiation from surface at 80 K

Thermal emission of aerosols

Thermal emission of the UV-absorbing aerosols explain the observations at 10 and 20 μm .

Titan's thermal emission (8-13 μm): a better resolution

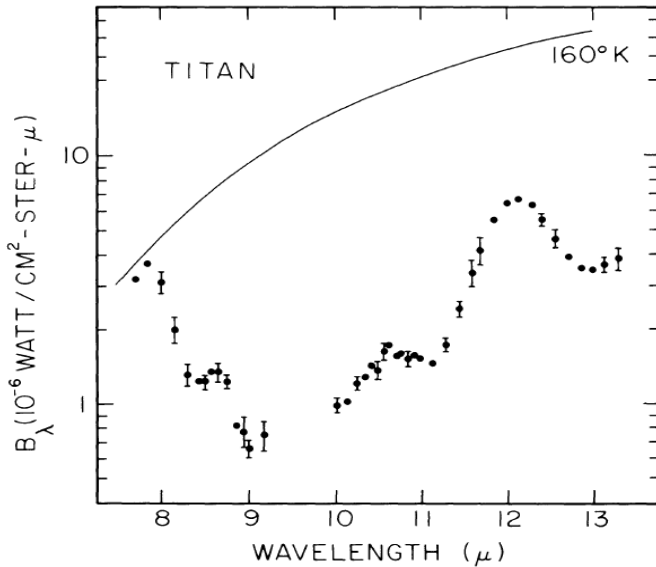


Gillett et al, 1975: observations at Kitt Peak (2.1 m and 4 m), assumed : $R_{\text{Titan}} = 2900 \text{ km}$.

They fitted the data with a model including a temperature inversion at 160 K (stratosphere)

- $q_{\text{C}_2\text{H}_6} = 0.5 \text{ atm-cm}$
- $q_{\text{C}_2\text{H}_2} = 1.0 \text{ atm-cm}$
- optically thin dust

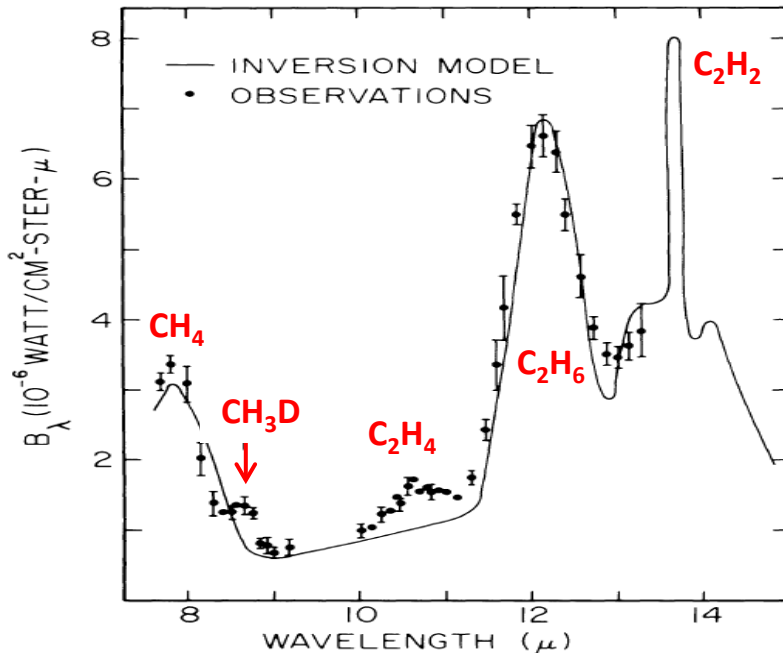
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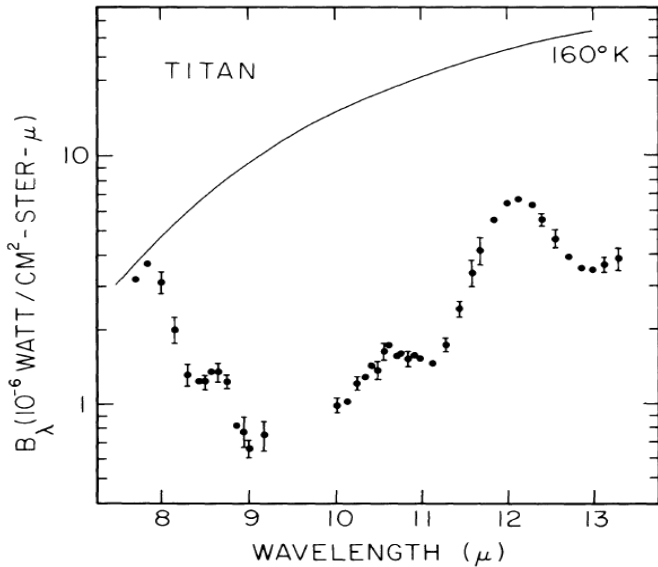
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- optically thin dust



They derived :

- $q_{\text{C}_2\text{H}_4} = 0.002 \text{ atm-cm}$ (in agreement with photochemical model)
- $q_{\text{CH}_3\text{D}} \sim 0.5 \text{ atm-cm}$

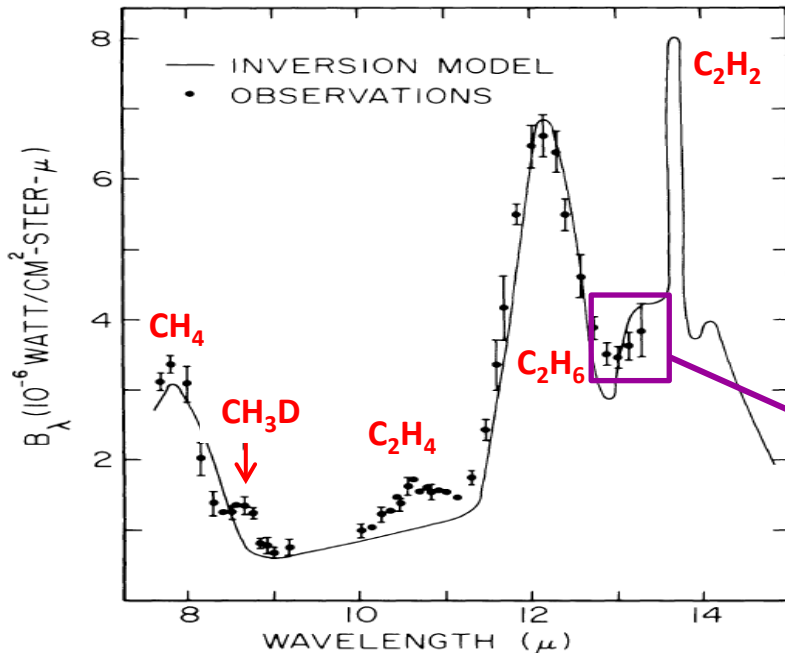
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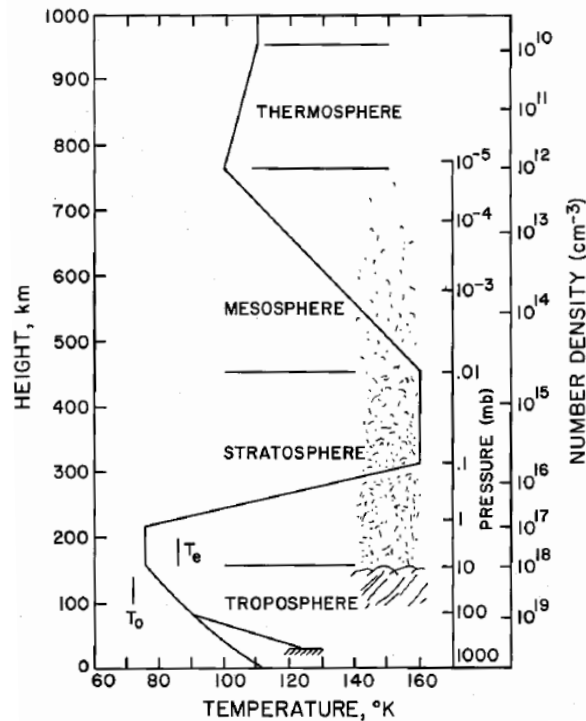
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Increase towards higher wavelength
-> suggest C_2H_2 , needed confirmation.

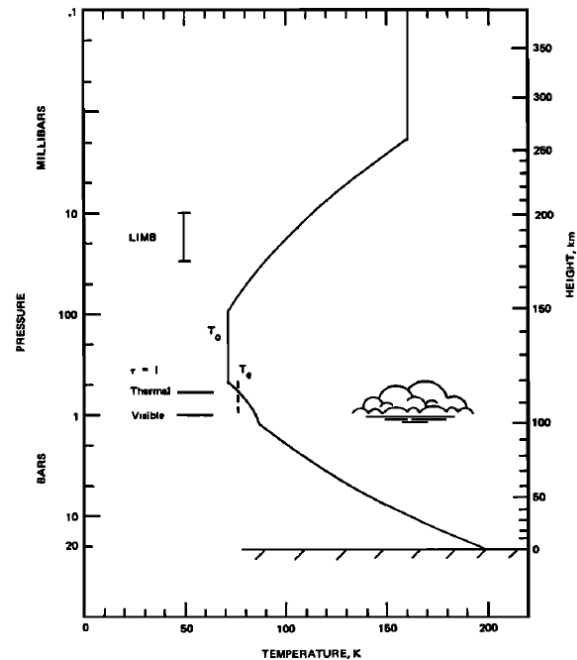
Titan's thermal profiles: models

- atmosphere of CH_4 in equilibrium with a solid CH_4 surface with $T_{\text{surf}} = 86 \text{ K}$ and $P_{\text{surf}} = 20 \text{ mbar}$ (Calwell, 1978).
- atmosphere of N_2 ($\text{CH}_4 = \text{minor constituent}$) with $T_{\text{surf}} = 200 \text{ K}$ and $P_{\text{surf}} = 20 \text{ bars}$ (Hunten, 1978) = greenhouse model due to pressure-induced absorption in N_2 .

Atmosphere of CH_4 , $m=16$



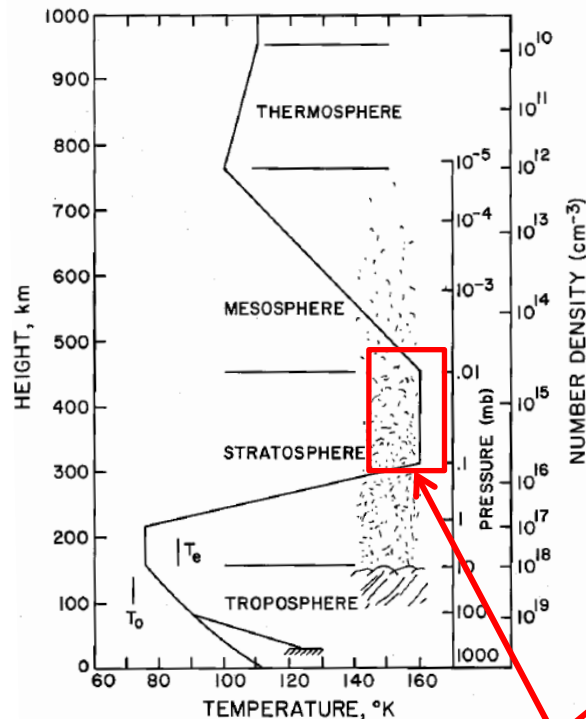
Atmosphere of N_2 , $m=28$



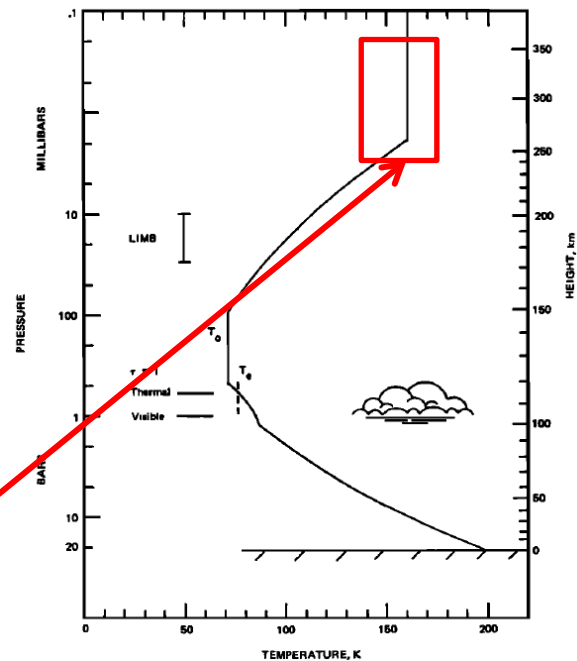
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Atmosphere of CH_4 , $m=16$



Atmosphere of N_2 , $m=28$



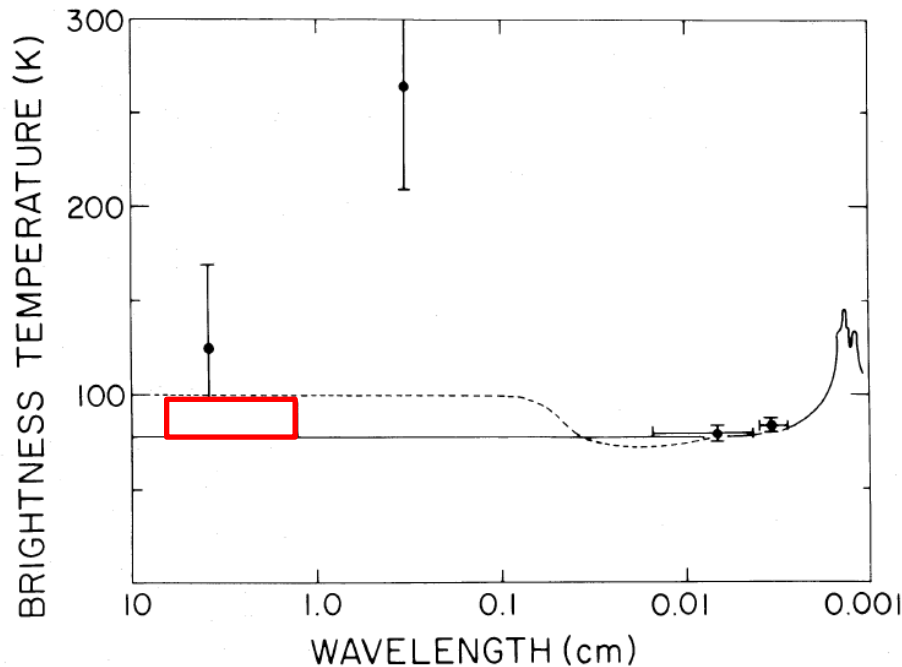
Consensus: T inversion with $\sim 160 \text{ K}$ (from the thermal emission of CH_4 at $7.7 \mu\text{m}$)

Titan in far IR (0.3- 6 cm): surface temperature

Radio wavelength -> no gas absorption in this spectral region, except NH_3 , but was not detected in Titan's atmosphere => **should probe the surface temperature**

- Jaffe et al., 1980 : observation with the VLA (interferometer) at 1.3, 2 and 6 cm.

-> derived $T_b = 87 \pm 9 \text{ K}$



Knowledge on Titan's atmosphere before space probes

- Molecular gas detected : CH_4 , C_2H_6 , C_2H_2 , CH_3D , C_2H_4 and maybe H_2
-> poor constrains of their column densities.
- Aerosols absorbing in UV => heating of the stratosphere = temperature inversion.
- Thermal emission of CH_4 at $7.7 \mu\text{m}$ => stratosphere temperature = 160 K.
- Cold surface temperature < 100 K.

Uncertainty on the main composition:

- atmosphere of CH_4
- atmosphere of N_2 (CH_4 = minor constituent)

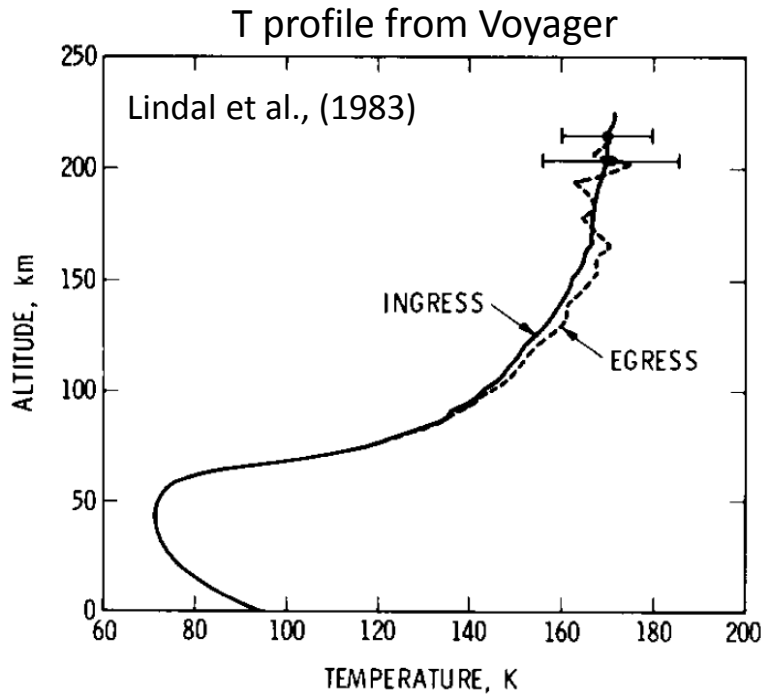
Some photochemical models => based on $\text{CH}_4 + \text{H}_2$ photochemistry

Titan's atmosphere derived from Voyager 1 & 2 (1980, 1981)

■ Thermal profile

Inferred from radio occultation measurements (Lindal et al., 1983)

- Radio occultation measurements => T/m profile
 - Mid-IR spectrometry at 540 cm^{-1} => $T_{\text{surf}} = 94\text{-}97\text{ K}$
- } $m \sim 28\text{ amu} \Rightarrow$ atmosphere of N_2 or CO

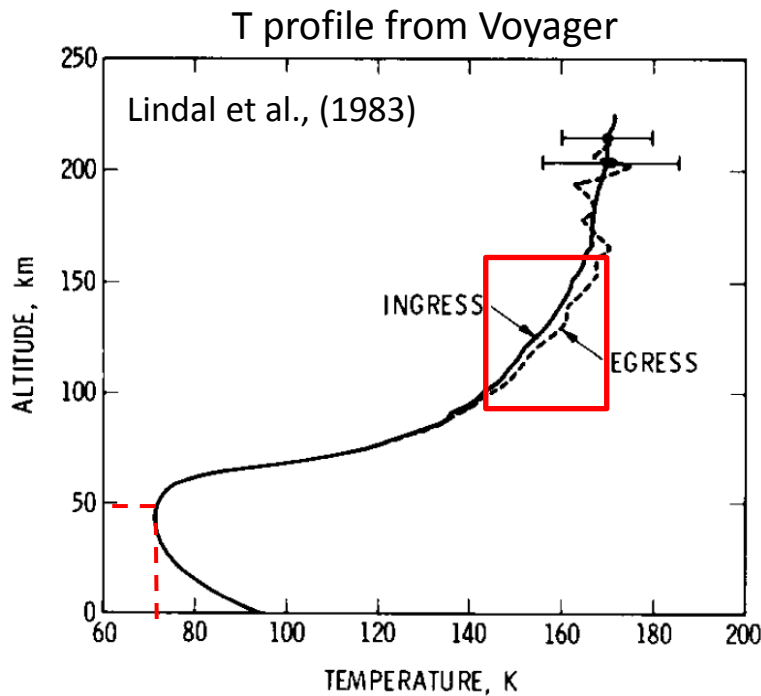


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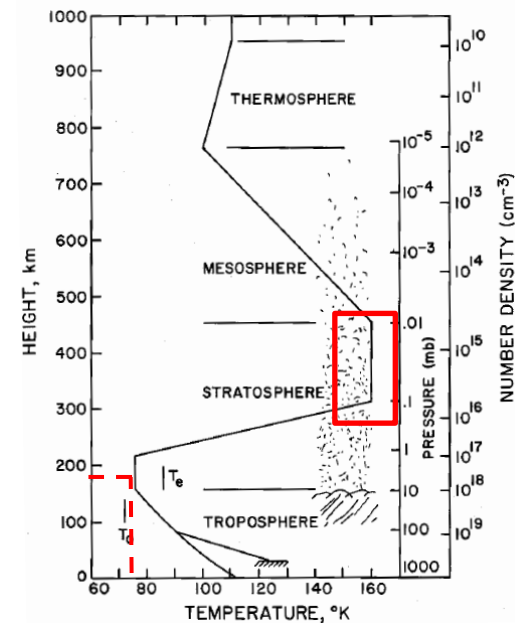
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Before Voyager (Hunten, 1978)



The predicted surface and tropopause temperature were close to the Voyager observations. The tropopause was thought to be 150 km higher than observed. Stratospheric temperature were correct.

Titan's atmosphere derived from Voyager 1 & 2 (1980, 1981)

■ Composition

- **UVS solar occultations and airglow measurements:** emission lines of N and N₂
 - => N₂ is the major component (82-95 %)
 - => CH₄ is a minor constituent (6% varying with altitude)

Titan's atmosphere derived from Voyager 1 & 2 (1980, 1981)

■ Composition

- **UVS solar occultations and airglow measurements:** emission lines of N and N₂
=> N₂ is the major component (82-95 %)
=> CH₄ is a minor constituent (6% varying with altitude)

- Thermal emission in mid-IR

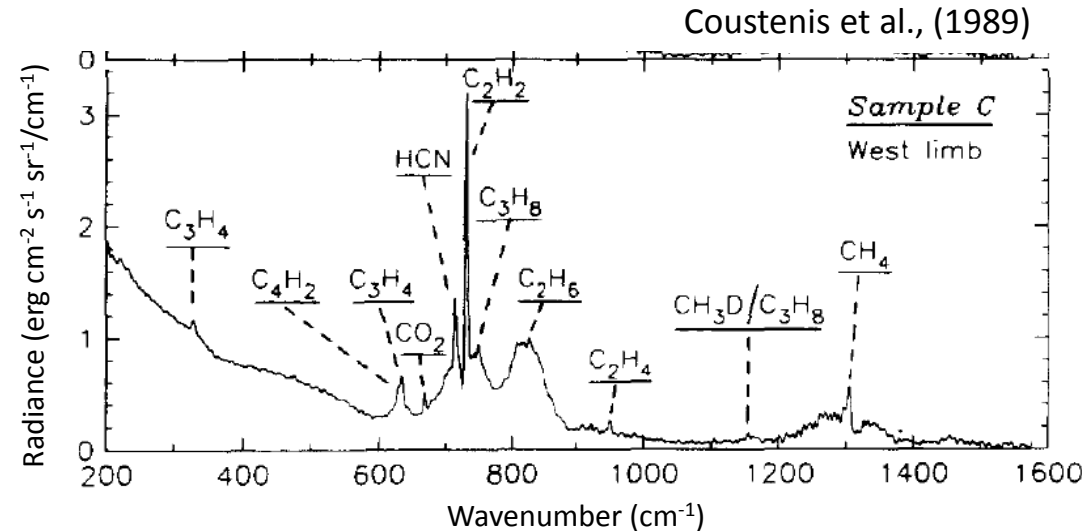
- In the deep atmosphere (troposphere) : **detection of N₂, CH₄ and H₂** (collision-induced absorption)

- In the stratosphere, detection of the emission bands of:

hydrocarbons: CH₄, CH₃D, C₂H₂, C₂H₆, C₂H₄, C₃H₈, CH₃C₂H, C₄H₂

nitriles: HCN, HC₃N, C₂N₂

oxygen compound: CO₂

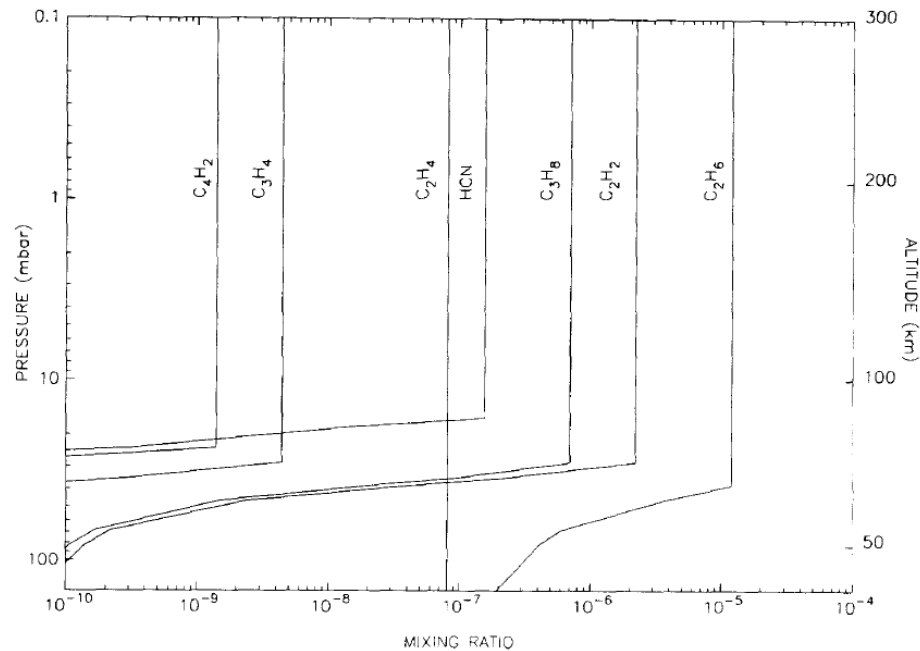


Titan's atmosphere derived from Voyager 1 & 2 (1980, 1981)

■ Composition

As the thermal profile was known => constrains of the molecular abundances with radiative transfer models.

But nadir observations => no vertical resolution



1995 : first General Circulation Model of Titan's atmosphere => explanation of the spatial distribution of molecules

After Voyager: ground-based observations

■ Observation in near IR

- 1983: CO detection at 4.5 μm , $q_{\text{CO}} = 6 \times 10^{-5}$ (Lutz et al 1983)

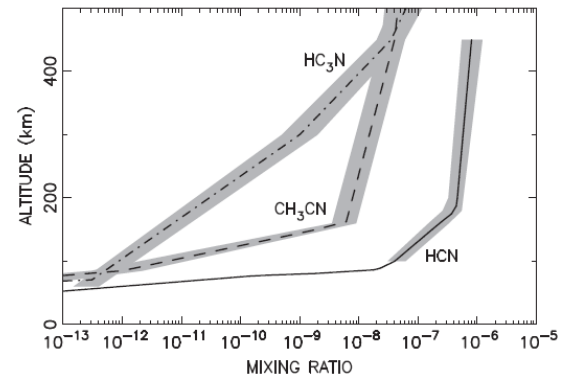
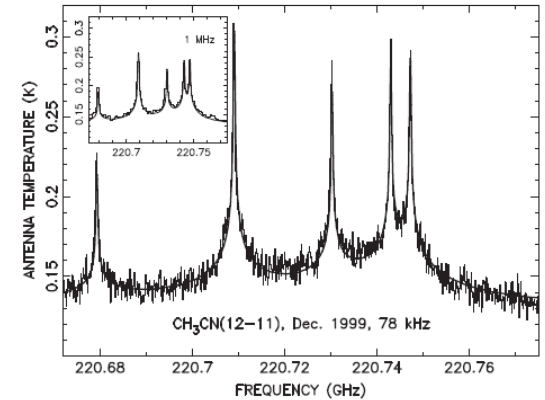
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■ Observations in mm/submm : very high resolution

- Detection of CH_3CN
- Vertical profiles of nitriles outside the Voyager range (from the line profile)
- $^{15}\text{N}/^{14}\text{N}$ in HCN, $^{18}\text{O}/^{16}\text{O}$ in CO
- Wind speed from the Doppler shifts of the lines .



After Voyager: ground-based observations

■ Observation in near IR

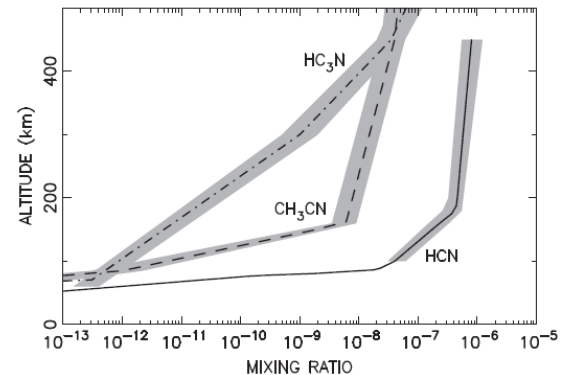
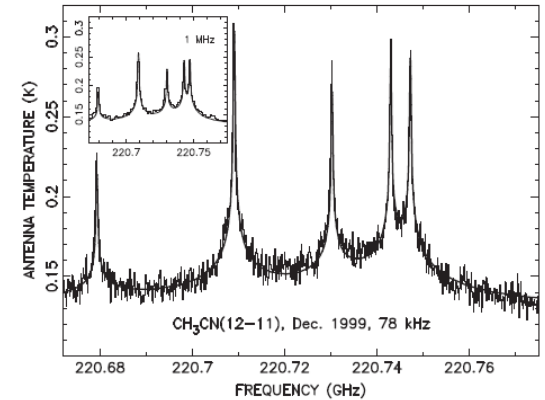
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■ Observations in infrared

- High resolution ground based observations:
 - > thermal profile from CH_4 lines at $7.7 \mu\text{m}$ in the 300-600 km range (higher than Voyager)
 - > wind velocity from Doppler shift from C_2H_6 lines at $12 \mu\text{m}$



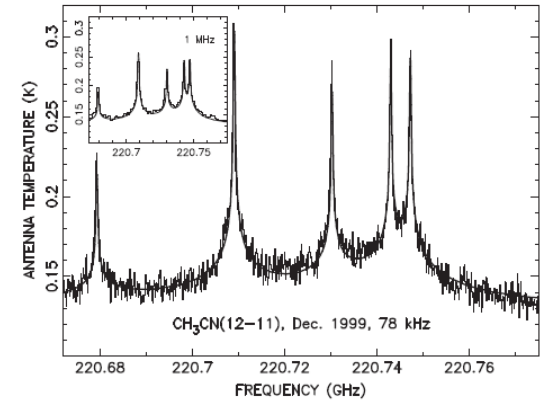
After Voyager: ground-based observations

■ Observation in near IR

- 1983: CO detection at $4.5 \mu\text{m}$, $q_{\text{CO}} = 6 \times 10^{-5}$ (Lutz et al 1983)

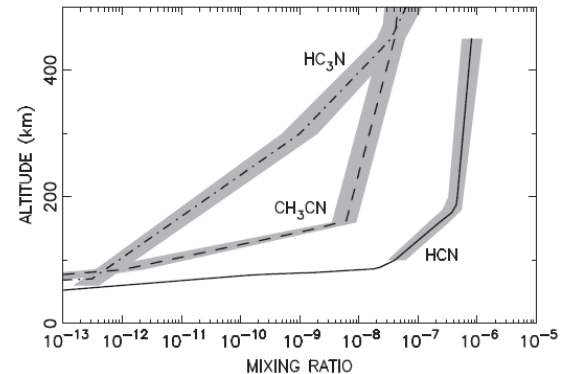
■ Observations in mm/submm : very high resolution

- Detection of CH_3CN
- Vertical profiles of nitriles outside the Voyager range (from the line profile)
- $^{15}\text{N}/^{14}\text{N}$ in HCN, $^{18}\text{O}/^{16}\text{O}$ in CO
- Wind speed from the Doppler shifts of the lines



■ Observations in infrared

- High resolution ground based observations:
 - > thermal profile from CH_4 lines at $7.7 \mu\text{m}$ in the 300-600 km range (higher than Voyager)
 - > wind velocity from Doppler shift from C_2H_6 lines at $12 \mu\text{m}$



■ Space observations (ISO)

detection of H_2O and C_6H_6 (Coustenis et al., 1998, 2003).

Titan's composition: before and after space probes

	Before Voyager	Voyager	Ground/satellites	Cassini
CH ₄	1 - 2 km. Am			
H ₂	< 2 km.Am			
C ₂ H ₂	1 cm. Am			
C ₂ H ₆	0.5 cm. Am			
C ₂ H ₄	2. 10 ⁻³ cm.Am			
C ₃ H ₈				
CH ₃ C ₂ H				
HCN				
HC ₃ N				
CH ₃ CN				
CO ₂				
CO				
H ₂ O				
C ₆ H ₆				

- the main component was unknown
⇒ impossible to infer abundances
- the level probed for each molecule is unknown:
column density ratio ≠ abundances ratio

Before Voyager : only column densities inferred for CH₄, H₂, C₂H₂, C₂H₆ and C₂H₄.

Coupling between T and composition -> **importance of the thermal profile observed with Voyager**

Titan's composition: before and after space probes

	"Before" Voyager ^(a)	Voyager	Ground/satellites	Cassini
CH ₄	1.3-2.6 %	6% (alt. var.)		4.9% surf, 1.48% atm.
H ₂	< 2.6 %	0.2 %		0.1 %
C ₂ H ₂	1 x 10 ⁻⁵	2.2 x 10 ⁻⁶	5.5 x 10 ⁻⁶ (ISO)	2.0 x 10 ⁻⁶
C ₂ H ₆	2 x 10 ⁻⁶	1.3 x 10 ⁻⁵	2.0 x 10 ⁻⁵ (ISO)	1.0 x 10 ⁻⁵
C ₂ H ₄	3 x 10 ⁻¹⁰	9 ± 5 x 10 ⁻⁸	1.2 x 10 ⁻⁷ (ISO)	1.0 x 10 ⁻⁷ (alt. var.)
C ₃ H ₈				
CH ₃ C ₂ H				
HCN				
HC ₃ N				
CH ₃ CN				
CO ₂				
CO				
H ₂ O				
C ₆ H ₆				

(a) Derived from column densities and Voyager results

Before Voyager : only column densities inferred for CH₄, H₂, C₂H₂, C₂H₆ and C₂H₄.
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After Voyager: **very good constraints because the T profile was known from Voyager**

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C ₂ H ₄	3 x 10 ⁻¹⁰	9 ± 5 x 10 ⁻⁸	1.2 x 10 ⁻⁷ (ISO)	1.0 x 10 ⁻⁷ (alt. var.)
C ₃ H ₈	-	7 ± 4 x 10 ⁻⁷	2 x 10 ⁻⁷ (ISO)	1.0 x 10 ⁻⁷
CH ₃ C ₂ H	-	4 ± 2 x 10 ⁻⁹	1.2 x 10 ⁻⁸ (ISO)	8 x 10 ⁻⁹
HCN	-	1.6 x 10 ⁻⁷	5 x 10 ⁻⁷ (alt. var.)	1 x 10 ⁻⁷
HC ₃ N	-	< 1.5 x 10 ⁻⁹	10 ⁻¹² -10 ⁻⁸ (alt.var.)	1 x 10 ⁻⁹
CH ₃ CN	-	-	1 x 10 ⁻⁸ (alt. var.)	-
CO ₂	-	1.4 x 10 ⁻⁸	2.0 x 10 ⁻⁸ (ISO)	1.6 x 10 ⁻⁸
CO	-	-	6 x 10 ⁻⁵ (near IR)	4.7 x 10 ⁻⁵
H ₂ O	-	-	8 ± 5 x 10 ⁻⁹ (ISO)	4 x 10 ⁻¹⁰
C ₆ H ₆	-	-	4 x 10 ⁻¹⁰ (ISO)	4 x 10 ⁻¹⁰

(a)Derived from column densities and Voyager results

Before Voyager : only column densities inferred for CH₄, H₂, C₂H₂, C₂H₆ and C₂H₄.

coupling between T and composition -> **importance of the thermal profile observed with Voyager** .

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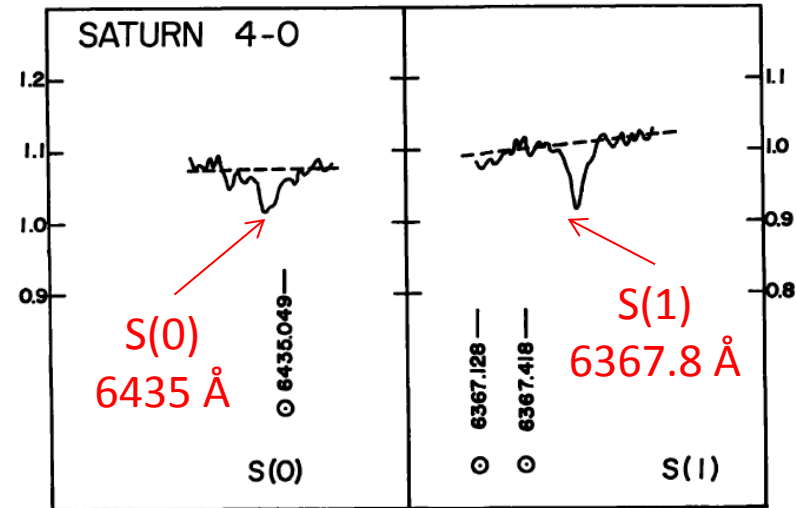
Saturn

Saturn's composition: before space probes flybys

- 1932: identification of CH_4 and NH_3
- 1962: first observation of H_2 from S(1) and S(0) lines in the (4-0) band (Spinrad et al., 1962)

From the line width:

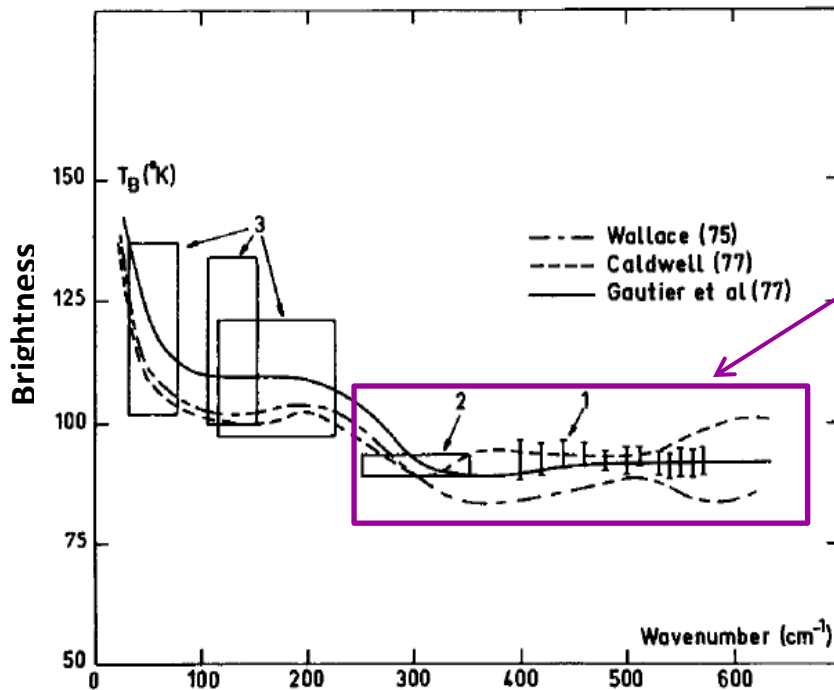
$$a_{\text{H}_2} = 76 \pm 20 \text{ km.Am (encrenaz 1973)}$$



H_2 detection by Encrenaz et al., 1973

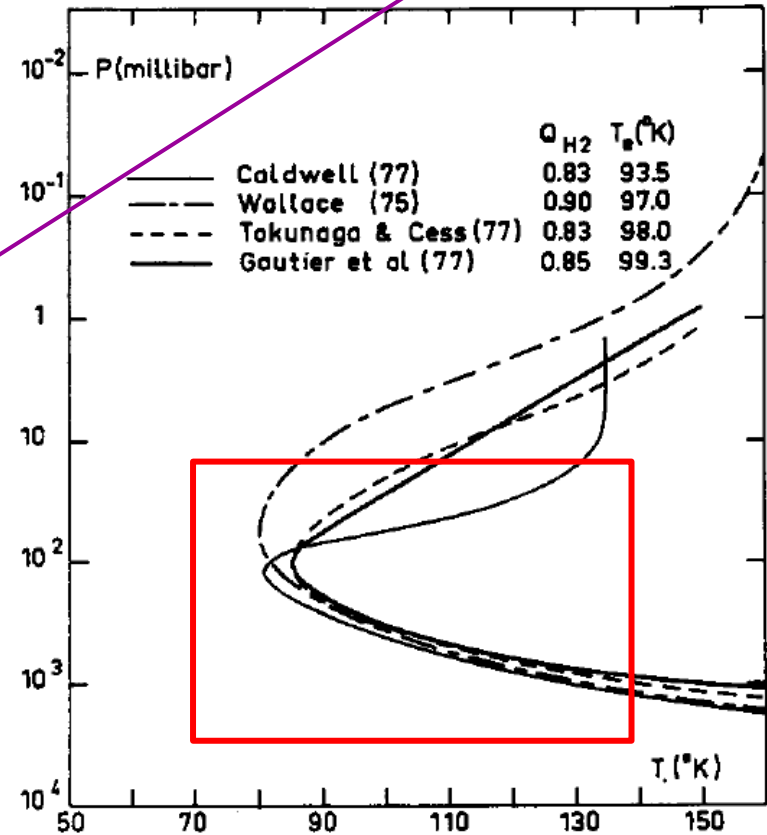
Saturn's thermal profile before space probes flybys

Fit of the H₂-H₂ and H₂-He collision-induced absorption between 250 and 800 cm⁻¹ (40-12 μm) => constraints of T in the deep atmosphere



Gautier et al. 1977

$q_{H_2} = 0.7, 0.85, 1.0$

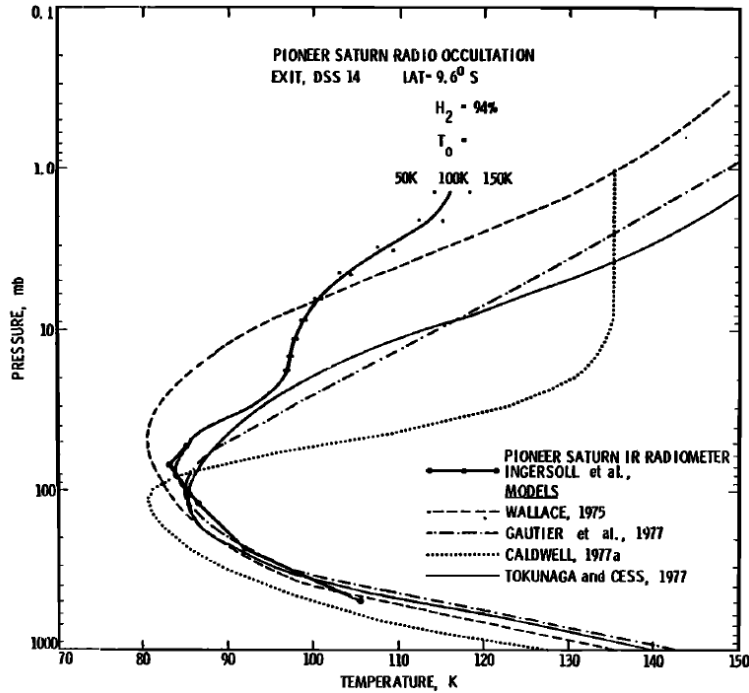


But the equilibrium temperature of Saturn should be 75 K (including ring shadowing)

=> Internal source of heating

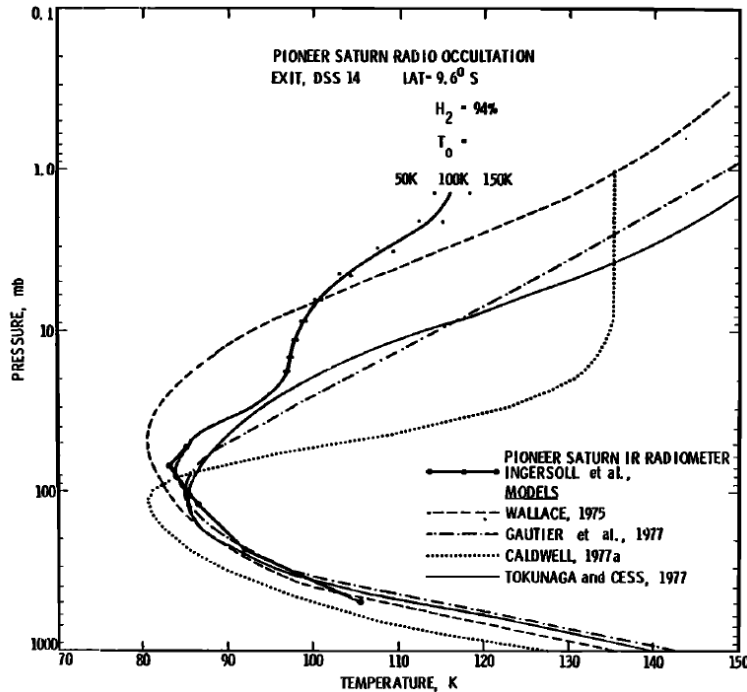
Saturn's temperature: comparison with probes results

Pioneer flyby (Sept. 1979)

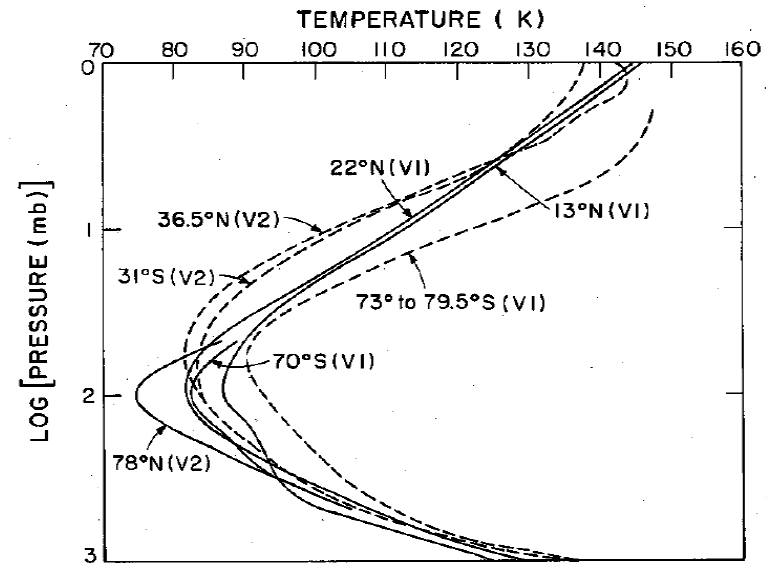


Saturn's temperature: comparison with probes results

Pioneer flyby (Sept. 1979)



Voyager 1 (V1) and Voyager 2 (V2) flybys

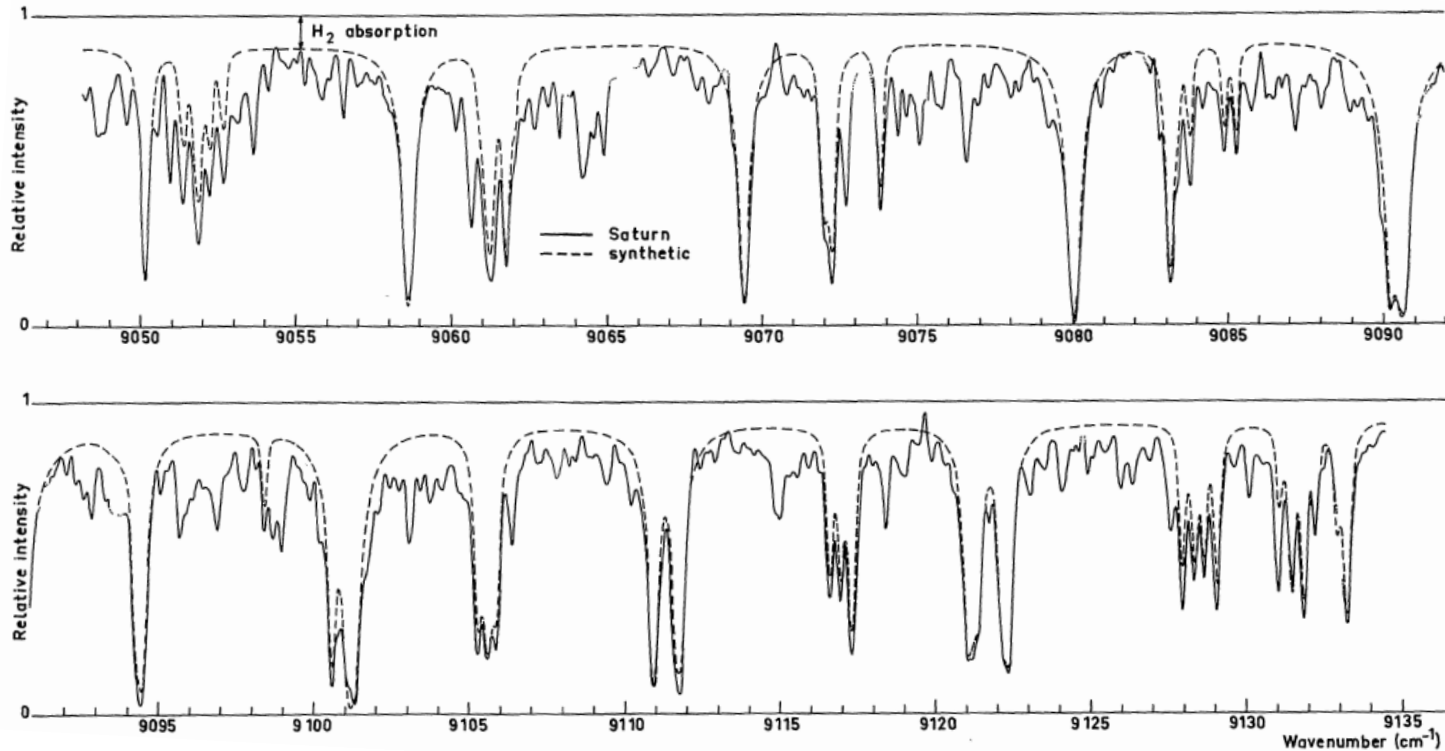


Good constraints on the deep thermal profile from H₂-H₂ and H₂-He collision-induced absorption spectrum

-> independent from the "insitu" spacecraft measurements

Saturn's composition: constraints before the probes

- **Case of CH₄ and H₂** from the 3ν₃ CH₄ band (0.83 - 1.67 μm) -> many CH₄ lines over a continuum due to H₂ pressure-induced absorption



Buriez and deBergh (1981)

$$\Rightarrow \text{CH}_4/\text{H}_2 = (4 \pm 2) \cdot 10^{-3}$$

From Voyager: $(2 - 4) \cdot 10^{-3}$

From Cassini: $(4.7 \pm 0.2) \cdot 10^{-3}$



H₂ and CH₄: Good constraints from ground-based observations

Saturn's composition: constraints before the probes

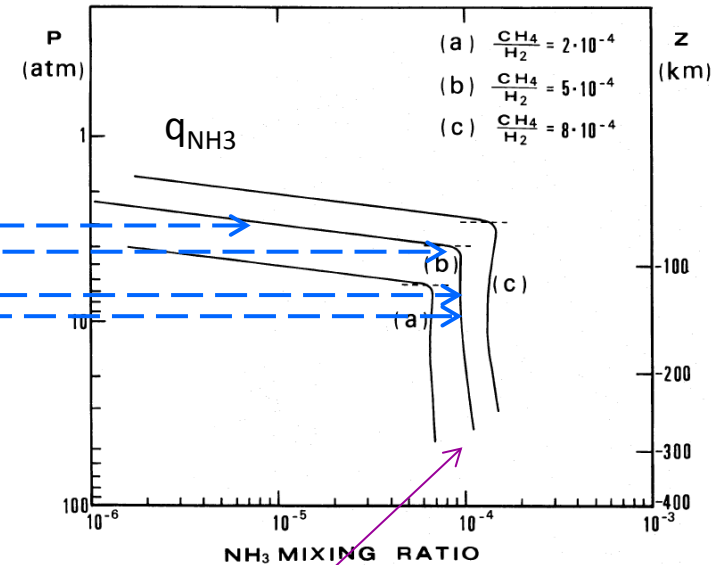
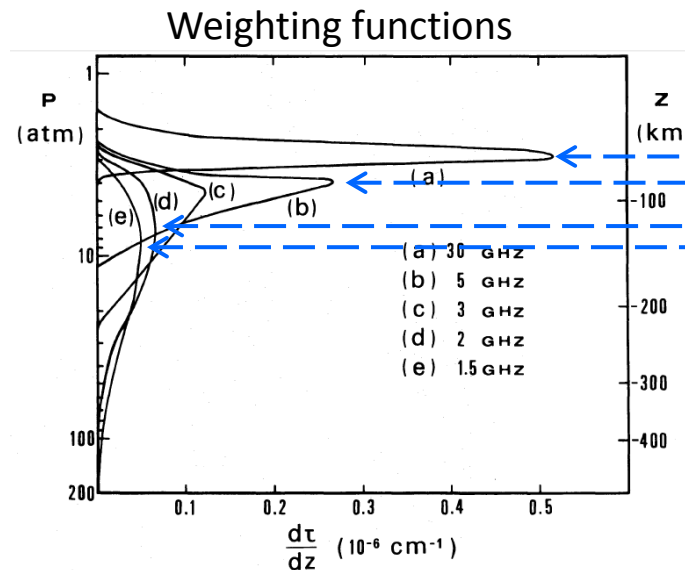
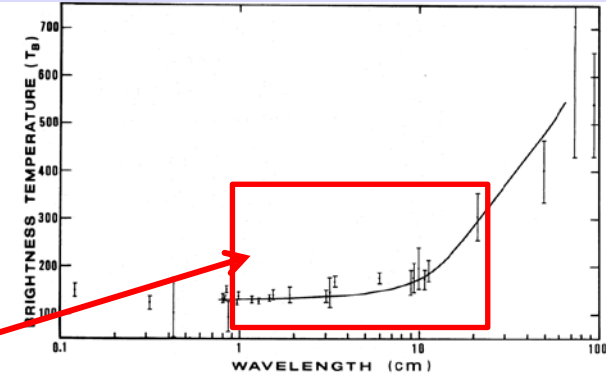
■ Case of NH_3

from its absorbing 1.25 cm band (Ohring et al., 1976)

No constraints of the tropospheric temperature

→ assumed adiabatic extrapolations, $\text{CH}_4/\text{H}_2 = 5 \cdot 10^{-4}$

They inverted the 1.5 – 30 cm spectral region



First vertical profile of NH_3 : $q = 1 \cdot 10^{-4}$ below the condensation level

From Voyager : $(0.5 - 2) \cdot 10^{-4}$, Courtin et al. (1984)

From Cassini : 1 - 3 bar : $q_{\text{NH}_3} = (1.4 - 5) \cdot 10^{-4}$ with latitudinal variations, Fletcher et al. (2011)

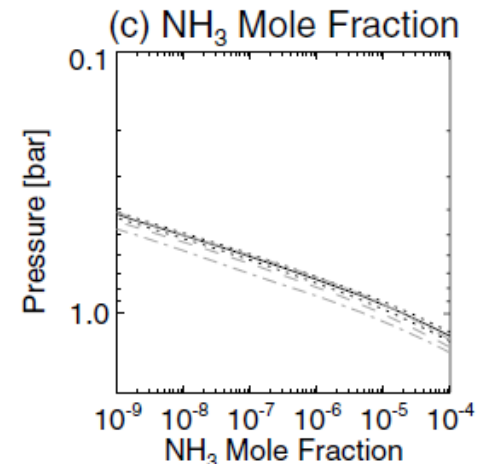
Saturn's composition: constraints before the probes

■ Case of NH₃

Many other constrains from the ground and with ISO and Herschel satellites

Reference	q_{NH_3}	Method
Courtin et al. (1984)	$(0.5 - 2.0) \times 10^{-4}$	Voyager/IRIS 180–300 cm ⁻¹
de Pater and Massie (1985)	5×10^{-4} at $p > 3$ bar 3×10^{-5} at $p < 1.25$ bar	Very Large Array (VLA)
Briggs and Sackett (1989)	$0.7 - 1.1 \times 10^{-4}$ at $p = 2$ bar	Radio T_B
Grossman et al. (1989)	1.2×10^{-4} around condensation level	VLA
Noll and Larson (1990)	Upper limit 3×10^{-4}	5 μm spectra
de Graauw et al. (1997)	1.1×10^{-4} at $p = 1.2$ bar	ISO/SWS
Kerola et al. (1997)	Less than 1×10^{-9} at radiative-convective boundary	3 μm data
Orton et al. (2000)	1×10^{-4} with 3–4 \times uncertainty	Sub-mm PH ₃ analysis
Burgdorf et al. (2004)	1×10^{-4}	ISO/LWS 96–101 cm ⁻¹
Kim et al. (2006)	6×10^{-8} at 460 mbar 3×10^{-8} at 390 mbar	3 μm data
Fletcher et al. (2009a)	$(3.3 \pm 0.3) \times 10^{-7}$ at 690 mbar	Cassini/CIRS far-IR

Vertical profile from the Herschel satellite : Fletcher et al., 2012



NH₃: good constraints from the ground-based and Earth satellites, independently of Saturn's probes.

Saturn's composition: hydrocarbons

■ C₂H₆

- From ground or Earth satellites :

$q_{\text{C}_2\text{H}_6} = 1.8 \times 10^{-6}$, 12.2 μm Tokunaga et al. (1975)

$(6 \pm 1) \times 10^{-6} < 20$ mbar, IUE (UV satellite), Winkelstein et al. (1983)

$(1.3 \pm 0.3) \times 10^{-5}$ at 0.5 mbar, ISO/SWS (infrared satellite), Moses et al. (2000)

$(1.5 \pm 0.5) \times 10^{-5}$ at 0.5 mbar, IRTF, Greathouse et al. (2005)

- From spacecrafts:

Voyager: $(3 \pm 1) \times 10^{-6} < 20$ mbar, Voyager/IRIS, Courtin et al. (1984)

Cassini: $(1.4 \pm 0.2) \times 10^{-5}$ at 1 mbar



C₂H₆: agreement between ground-based and spacecraft measurements

Saturn's composition: hydrocarbons

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😊 C₂H₆: agreement between ground-based and spacecraft measurements

■ Other hydrocarbons:

😊 C₂H₂: many measurements with IUE, TEXES/IRTF, SWS/ISO, Celeste/IRTF

😞 C₃H₈: $(2.6 \pm 0.8) \times 10^{-6}$ with TEXES/IRTF while $(0.9 - 1.5) \times 10^{-7}$ from Cassini/CIRS

😊 CH₃C₂H: 1 measurement with SWS/ISO , in agreement with Cassini/CIRS


😊 C₄H₂: 1 measurement with SWS/ISO , in agreement with Cassini/CIRS


C₆H₆: only detection with SWS/ISO, no detection from probes

Saturn's composition

■ Oxygen compounds

- **CO** : detection in submm but not with the probes

 - **CO₂** : 1 measurement from ISO/SWS (3.0×10^{-10} at 0.3-10 mbar), agrees with Cassini/CIRS

 - **H₂O** :
1.5 x 10⁻⁷ at 2 mbar from IUE (UV)

6.0 x 10⁻⁹ at 2 mbar from SWS/ISO (2.4 – 45 μm)

H₂O was not detected by Saturn's probes (because of bad S/N ratio)

Saturn's composition

■ Oxygen compounds

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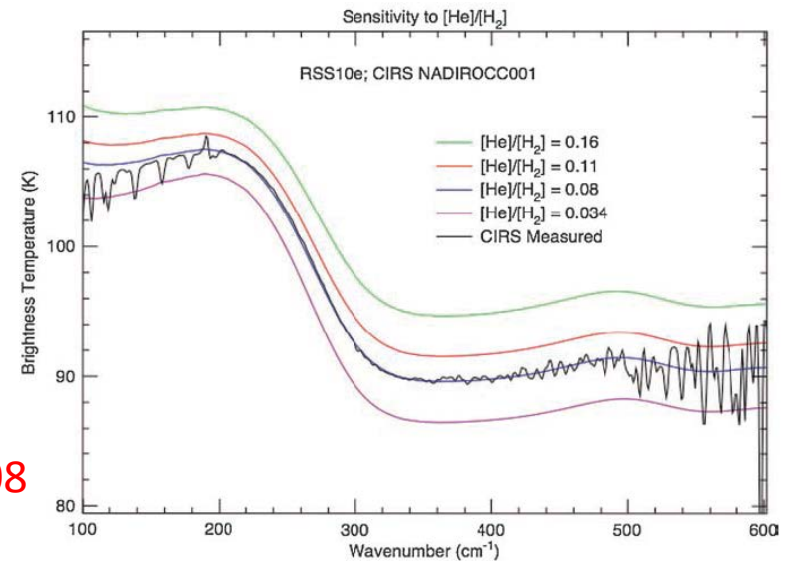
😞 - **H₂O** :
1.5 $\times 10^{-7}$ at 2 mbar from IUE (UV)
6.0 $\times 10^{-9}$ at 2 mbar from SWS/ISO (2.4 – 45 μm)

H₂O was not detected by Saturn's probes (because of bad S/N ratio)

■ Helium

Determined from Voyager and Cassini, impossible from the ground because of bad S/N ratio

From Cassini/CIRS: $[\text{He}]/[\text{H}_2]=0.08$



Jupiter

Jupiter knowledge before probes flybys

1932: detection of **CH₄** and **NH₃** (Wildt, 1932)

1960: detection of **H₂** (Kiess et al.)

Helium was presumed because of the theory of the primitive solar nebula composition

Jupiter knowledge before probes flybys

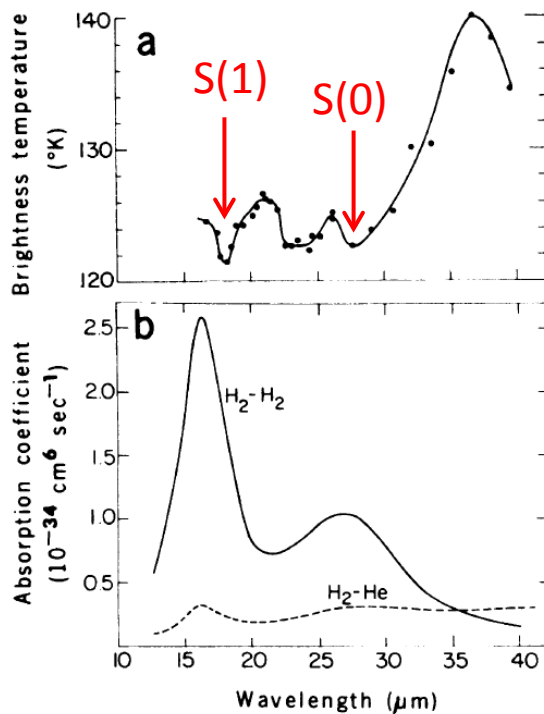
1932: detection of CH_4 and NH_3 (Wildt, 1932)

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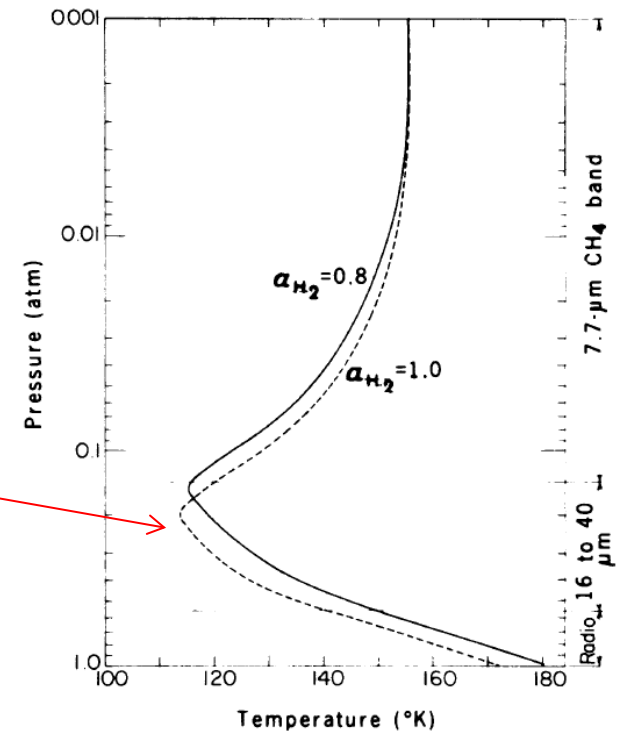
■ Thermal profile

constrained from H_2 - H_2 and H_2 -He collision induced absorptions (Houck et al., 1975)



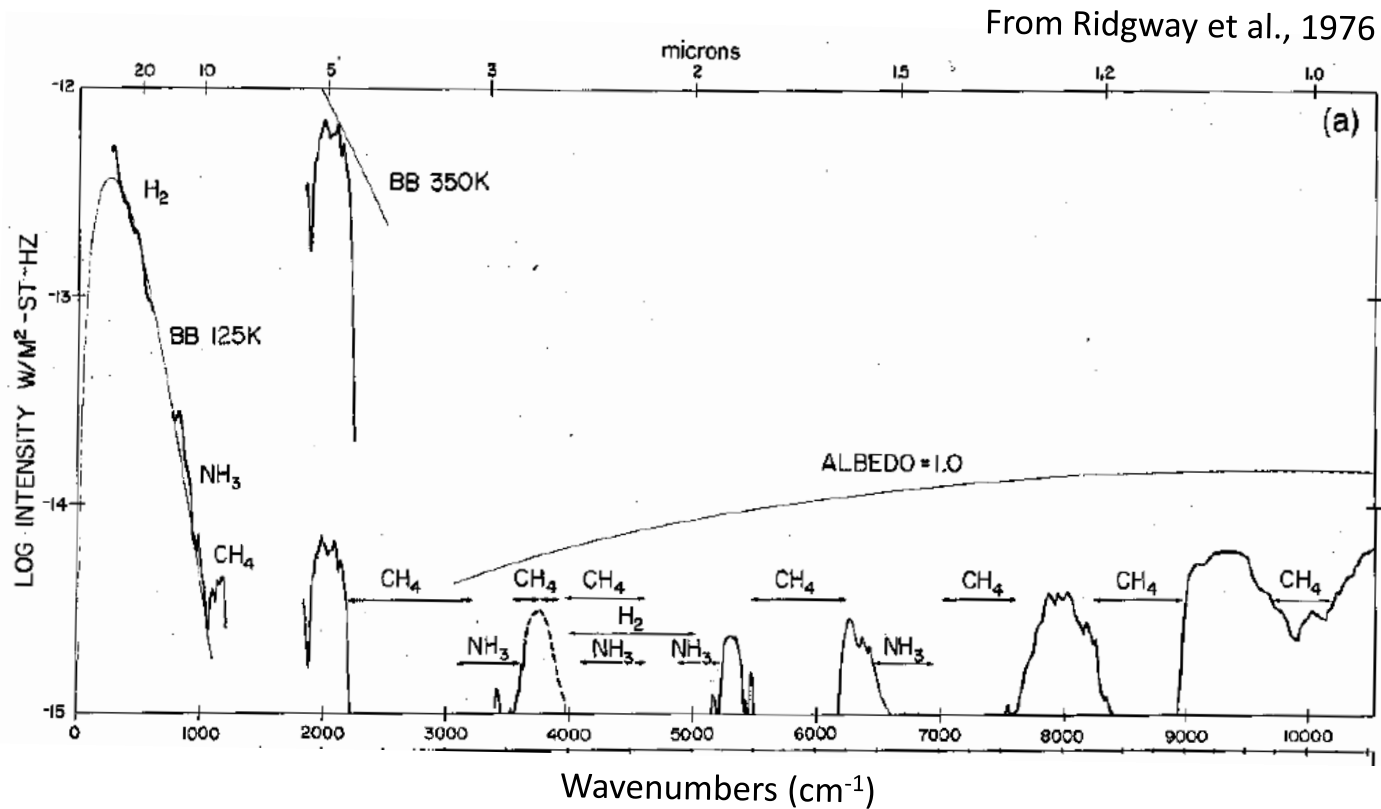
Results:

- $\alpha_{\text{H}_2} = 0.89 \pm 0.11$
- tropospheric temperature



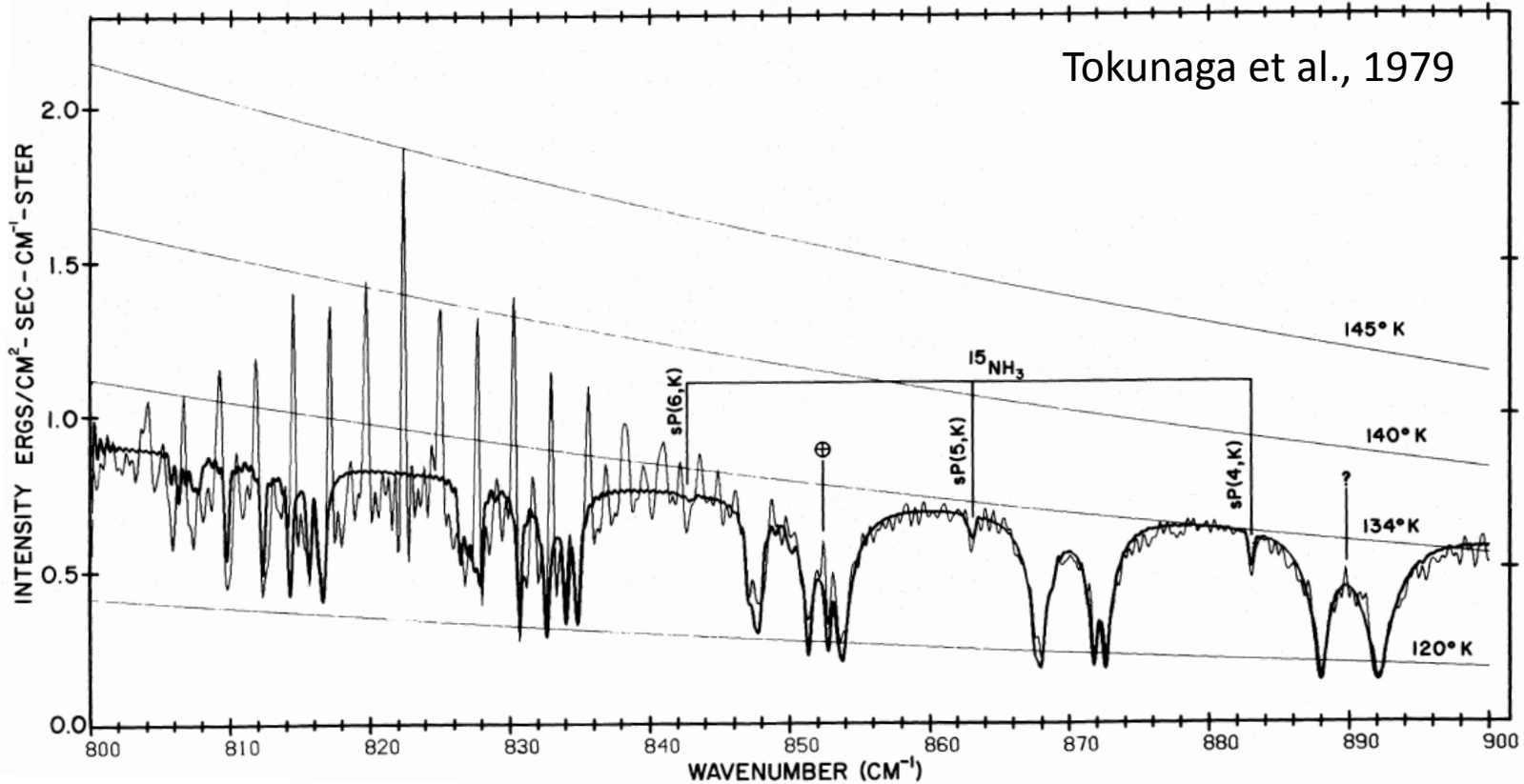
Jupiter knowledge before probes flybys

1975 : Jupiter's infrared spectrum



Jupiter composition before probes flybys

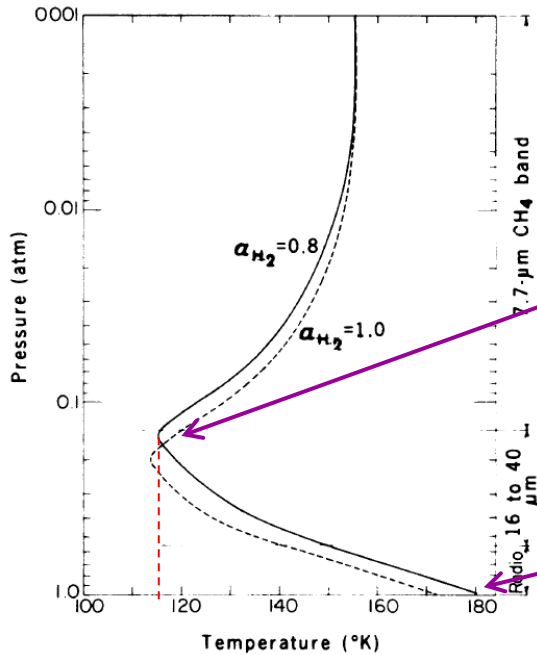
High spectral resolution observations in IR



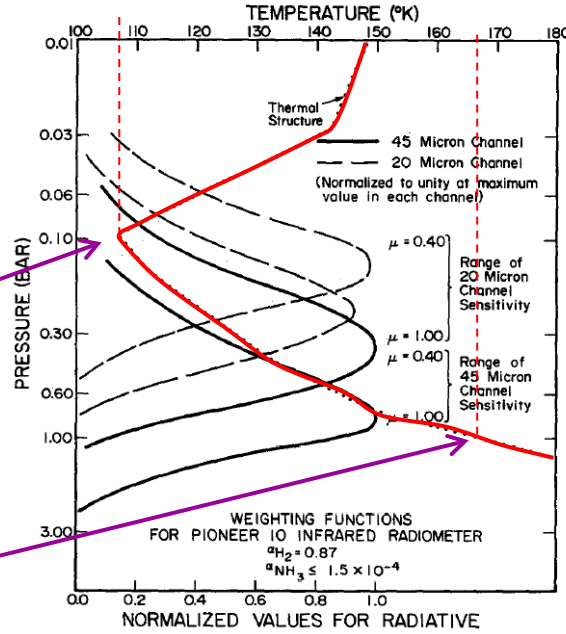
- line profile of NH₃ absorption lines => sensitivity to different levels , while continuum is sensitive to the NH₃ cloud deck
- C₂H₆ observed in emission => emission comes from a region where T ↗ with height

Jupiter thermal profile: comparison with probes

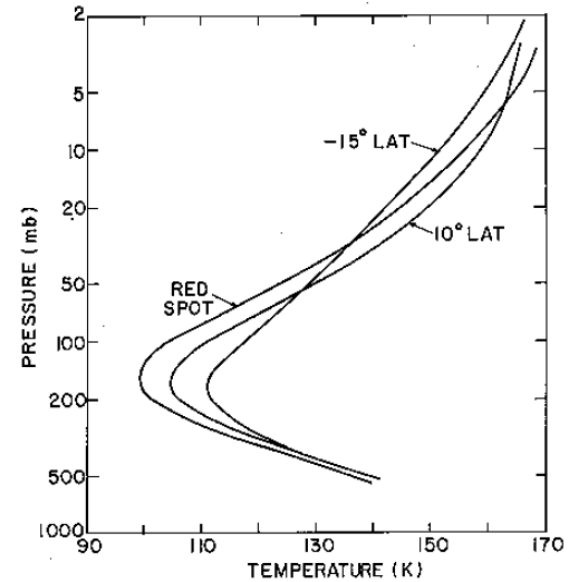
Before Pioneer, 1975



From Pioneer, Orton 1975












From Voyager, Hanel 1979



Good agreement from ground-based observations and Pioneer spacecraft

Jupiter composition: comparison with probes

■ Abundances, some examples

	Before spacecrafts (Ridgway, 1976)	From spacecrafts (Taylor, 2004)
 H ₂	0.89	0.86
 He	0.05 – 0.15	0.136
 CH ₄	7×10^{-4}	1.8×10^{-3}
 NH ₃	2×10^{-4}	7.0×10^{-4}
 H ₂ O	1×10^{-6}	$> 5.0 \times 10^{-4}$
Non equilibrium gas		
 CO	2×10^{-9}	1.5×10^{-9}
 PH ₃	4×10^{-7}	5×10^{-7}
Photochemical products		
 C ₂ H ₆	4×10^{-4}	3×10^{-6}
 C ₂ H ₂	8×10^{-5}	2×10^{-8}

Disagreement mostly come from the thermal profiles

Conclusions regarding the solar system observations

■ Giant planets

Main composition = H_2 (~ 90%) and He (~ 10 %).

T in the troposphere can be constrained by the absorption-induced bands of H_2 = most reliable probe for temperature sounding for giant planets.

⇒ Good estimation of the molecular abundances in the troposphere

⇒ But for photochemical product emitting in the stratosphere, constrains are poorer

-> need of high spectral resolution observations to probe higher.

Conclusions regarding the solar system observations

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⇒ Good estimation of the molecular abundances in the troposphere

⇒ But for photochemical product emitting in the stratosphere, constrains are poorer

-> need of high spectral resolution observations to probe higher.

■ Telluric planets

Very difficult ... because the main composition is a priori unknown.

Large spectral coverage => constrains on the T at several level: e.g. surface temperature in radio and stratospheric temperature in mid-IR.

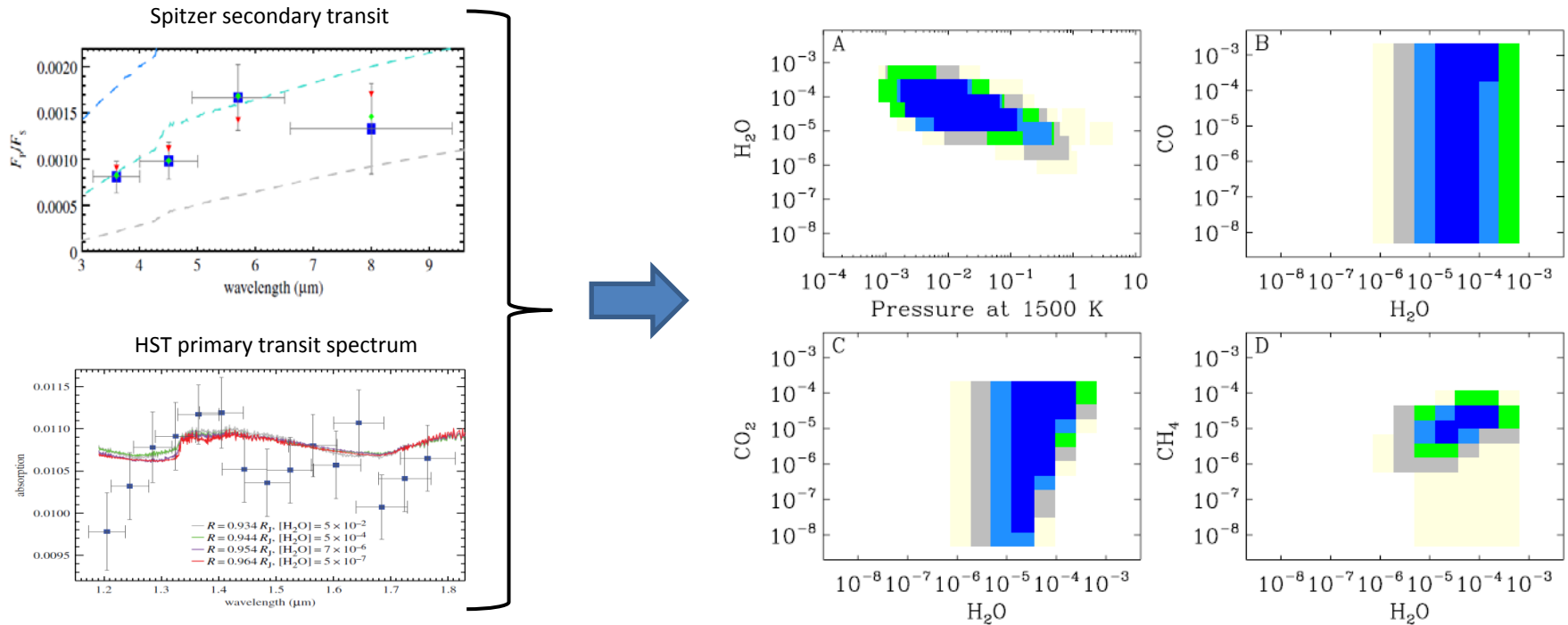
The knowledge of the T profile from Voyager observations was crucial in the knowledge of Titan's composition.

Perspectives: 2 promising methods for hot Jupiters

■ Use of primary transit and secondary transit spectra simultaneously (Griffith et al., 2014)

- **Primary transit** -> terminator, absorption of the stellar light => sensitive to the planet composition and not so much to the T profile
- **Secondary transit** -> day side, stellar light diffusion + thermal emission => sensitive to the composition and the T profile

Fit of the secondary transit + primary transit of XO-2b (include H₂O, CH₄, CO, CO₂)



Need of large spectral range and high resolution to decouple T and composition.

Perspectives: 2 promising methods for Hot Jupiters

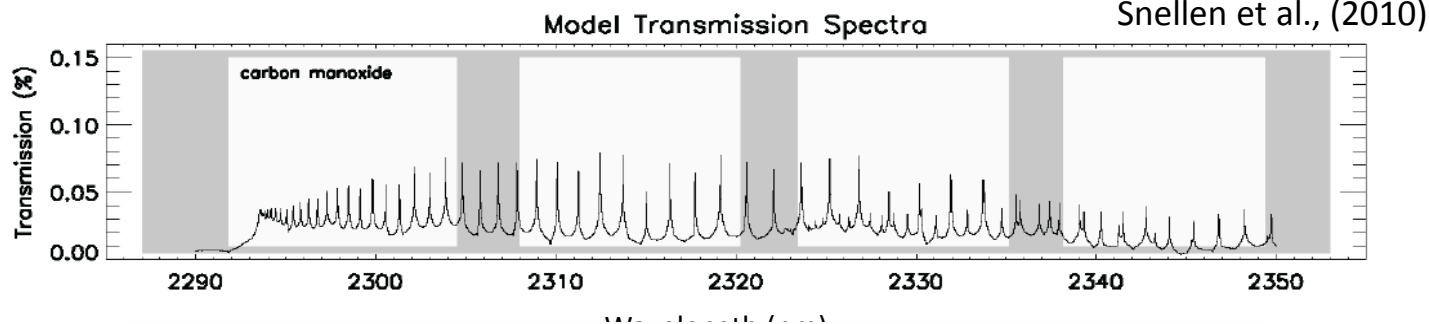
■ Detection of a molecule from high resolution spectra (Snellen et al. 2010)

High spectral resolution observation of exoplanets with VLT/CRILES ($R = 100\,000$)

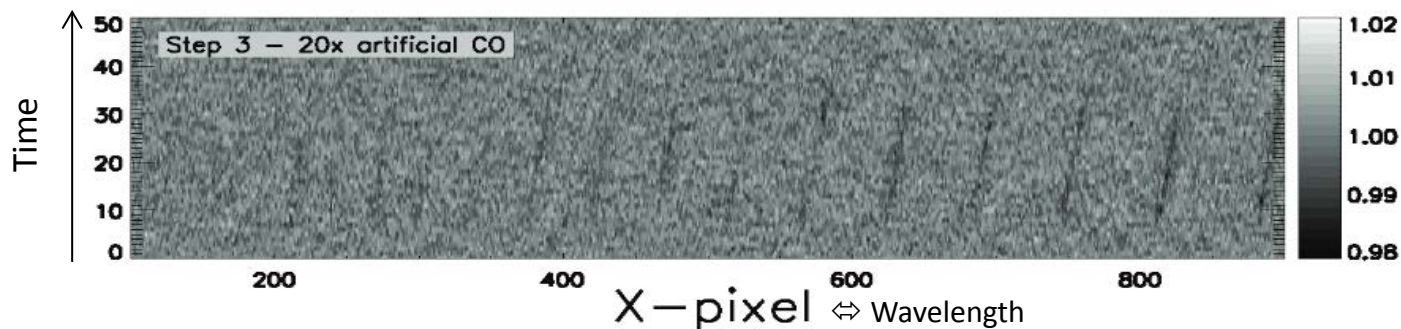
Principle: molecular lines are Doppler shifted while the planet is moving on its orbit, whereas the stellar spectrum is not.

=> Case of CO in HD209458b, observation during the primary transit

Model of the CO band at $2.3\ \mu\text{m}$



Doppler shifted lines during the transit



Detection of CO in transit: HD209 and HD189
outside transit: τ -Boo and 51 Peg